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1	Neural variational Data Assimilation with Uncertainty Quantification using
2	SPDE priors
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ABSTRACT: The spatio-temporal interpolation of large geophysical datasets has historically been 11 addressed by Optimal Interpolation (OI) and more sophisticated equation-based or data-driven 12 Data Assimilation (DA) techniques. Recent advances in the deep learning community enables 13 to address the interpolation problem through a neural architecture incorporating a variational 14 data assimilation framework. The reconstruction task is seen as a joint learning problem of 15 the prior involved in the variational inner cost, seen here as a projection operator of the state, 16 and the gradient-based minimization of the latter. Both prior models and solvers are stated 17 as neural networks with automatic differentiation which can be trained by minimizing a loss 18 function, typically the mean squared error between some ground truth and the reconstruction. 19 Such a strategy turns out to be very efficient to improve the mean state estimation, but still needs 20 complementary developments to quantify its related uncertainty. In this work, we use the theory 21 of Stochastic Partial Differential Equations (SPDE) and Gaussian Processes (GP) to estimate both 22 space-and time-varying covariance of the state. Our neural variational scheme is modified to 23 embed an augmented state formulation with both state and SPDE parametrization to estimate. We 24 demonstrate the potential of the proposed framework on a spatio-temporal GP driven by diffusion-25 based anisotropies and on realistic Sea Surface Height (SSH) datasets. We show how our solution 26 reaches the OI baseline in the Gaussian case. For nonlinear dynamics, as almost always stated 27 in DA, our solution outperforms OI, while allowing for fast and interpretable online parameter 28 estimation. 29

30 1. Introduction

Over the last decade, the emergence of large spatio-temporal datasets both coming from remote 31 sensing satellites and equation-based numerical simulations has been noticed in Geosciences. As 32 a consequence, the need for statistical methods able to handle both the size and the underlying 33 physics of these data is growing. Data assimilation (DA) is the traditional framework used by 34 geoscientists to merge these two types of information, data and model, by propagating information 35 from observational data to areas with missing data. When no deterministic numerical outputs 36 are available, the classic approach stems from the family of the so-called Optimal Interpolation 37 (OI) techniques, also being at the core of the statistical DA methods (Asch et al. 2016). For the 38 production of gridded geophysical maps, most of operational products either derived from OI 39 and DA schemes allow us to estimate the mesoscale components of the targeted variables. In 40 the specific case of Sea Surface Height (SSH) for instance, even if satellite altimetry provides 41 capabilities to inform the mesoscale ocean geostrophic currents, the gridded product fails in 42 merging observations and background model with consistent temporal and spatial resolutions able 43 to retrieve fine mesoscale structures lower than 150-200 km at mid-latitudes, while they are key 44 for general ocean circulation (Su et al. 2018). Then, recent efforts of the DA community have been 45 made to counteract this lack in the gridded SSH products, amongst them the dynamical optimal 46 interpolation (DOI) (Ubelmann et al. 2015), Multiscale Interpolation Ocean Science Topography 47 (MIOST) (Ubelmann et al. 2021) or BFN-QG (Le Guillou et al. 2023). 48

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From another point of view, deep learning frameworks are currently knowing an intense 50 period of scientific contributions to revisit statistical methods with neural network formulation. 51 State-of-the-art methods leverage DL to better extract the information from the observations 52 compared to the classical DA and OI approaches, in which fine-tuning of the DA scheme 53 parametrizations (background and observation error covariances) demands itself some method-54 ological and computational effort, see e.g. Tandeo et al. (2020), all together with the popular 55 Gaussian simplifications involved in most of DA operational schemes (Asch et al. 2016), though 56 some approaches exist to alleviate the Gaussian assumptions in nonlinear geophysical problems, 57 see e.g. Kurosawa and Poterjoy (2023). End-to-end learning architectures are also designed being 58 backboned on DA schemes (Boudier et al. 2023; Rozet and Louppe 2023; Fablet et al. 2021), so 59

that they draw from the bayesian formalism to learn all the components of the DA procedure (prior model, observation operator, numerical solver, etc.) at once. Last, over the last few years, the successfull applications of generative models to geophysical datasets paves the way to combine DL together with UQ by efficient sampling strategies of the posterior distribution. Though, to our knowledge, no end-to-end-combination of DA, DL and UQ has yet been proposed.

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That is why in this work, we draw from preliminary neural schemes inspired by variational data 66 assimilation, see e.g. (Fablet et al. 2021), to jointly learn the SPDE parametrization of a surrogate 67 stochastic prior model of the evolution equation, together with the solver of the minimization 68 problem. Because the parameters of the SPDE remain initially unknown, they are embedded in the 69 optimization process using an augmented state formulation, commonly used in data assimilation, 70 see e.g. (Ruiz et al. 2013). The solver is still based on an iterative residual scheme (Fablet et al. 71 2021) to update the analysis state. Here, the analysis stands for the expectation of the state given the 72 observations, together with the SPDE parametrization maximizing their likelihood given the true 73 states used during the training. The SPDE equation can then be seen as a tangent linear model of the 74 prior along the data assimilation window, from which we provide uncertainty quantification (UQ) 75 throughout its precision matrix. Using such a stochastic prior entails the possibility of generating 76 huge members in the posterior pdf, after conditioning of the prior samples by our neural variational 77 scheme. Also, if the training dataset is large enough, the method provides an efficient way to 78 estimate *online* the SPDE parametrization for any new sequence of input observations, without any 79 additional inference to make. In the end, the key contributions of this work are four-fold: 80

• We develop the explicit solver of the considered SPDE prior. It relies on the analytical expression for the SPDE-based precision matrix of any state trajectory, based on a finitedifference discretization of the grid covered by the tensors involved in our neural scheme;

We exploit this SPDE parametrization as surrogate prior model in the proposed variational
 formulation and leverage a trainable gradient-based solver to address jointly the interpolation
 of the state trajectory and the estimation of SPDE parameters from irregularly-sampled observations. The end-to-end training of the solver targets both the expectation of the state given
 the observations, together with the SPDE parametrization maximizing its likelihood given the
 true states;

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• The SPDE prior paves the way to uncertainty quantification through the sampling of the prior pdf and the conditioning by the neural gradient-based solver;

Two applications of this framework are provided: first, a GP driven by a spatio-temporal
 SPDE with spatially varying diffusion parameters is compared to the optimal solution; then
 a Sea Surface Height realistic dataset is used to demonstrate how the proposed framework is
 also relevant for non-gaussian and non-linear dynamics.

To present these contributions, the paper is structured as follows. In Section 2, we provide a preliminary background on SPDE-based Optimal Interpolation, Data assimilation and Deep Learning for DA, that we aims to bridge together in this work. We also remind how UQ is usually handled in this different fields. In Section 3, we present how we extend neural solvers to embed SPDE prior parametrizations for UQ. Finally, Section 4 provides two applications of this framework in the Gaussian case and for realistic Sea Surface Height datasets.

¹⁰² 2. Background: Data Assimilation, Uncertainty Quantification and Machine Learning

In this work, we target the reconstruction of a probabilistic spatio-temporal state sequence 103 $\mathbf{x} = {\mathbf{x}_k(\mathcal{D})}, \mathbf{x}_k \in \mathbb{R}^m$ given the partial and potentially noisy observational dataset $\mathbf{y}(\Omega) =$ 104 $\{\mathbf{y}_k(\Omega_k)\}, \mathbf{y}_k \in \mathbb{R}^{p_k}$, where $\Omega = \{\Omega_k\} \subset \mathcal{D}$ denotes the subdomain with observations and index 105 k refers to time t_k . To do so, we aim at bridging Data Assimilation (DA), Uncertainty Quantifi-106 cation (UQ) and deep learning-based (DL) methods, see Fig. 1 for a conceptual illustration of 107 these connections, to propose SPDE-based extensions of the so-called 4DvarNet neural variational 108 scheme (Fablet et al. 2021), as a generic interpolation and short-term forecasting tools. We provide 109 here a brief presentation of these three literatures, focusing only on what is useful in our framework, 110 with an additional description of how UQ is usually handled among these communities. For a 111 more exhaustive review, Cheng et al. (2023) provides a detailed presentation of machine learning 112 techniques with data assimilation and uncertainty quantification for dynamical systems. 113

a. Optimal Interpolation (OI) and Data Assimilation (DA)

(*i*) *Classic formulations*. As very basic details to ease the link with the other components of this work, we remind that when no dynamical model is available as a prior information, the

¹¹⁷ covariance-based Optimal Interpolation, see e.g. (Chilès and Delfiner 2012), is given by:

$$\mathbf{x}^{\star} = \mathbf{P}_{\mathbf{x}\mathbf{y}}\mathbf{P}_{\mathbf{y}\mathbf{y}}^{-1}\mathbf{y} \tag{1}$$

where \mathbf{P}_{xy} and \mathbf{P}_{yy} are covariance matrices coming from the covariance $\tilde{\mathbf{P}}$ of the observation and state vector $\begin{bmatrix} y & x \end{bmatrix}^{\mathrm{T}}$:

$$\tilde{\mathbf{P}} = \begin{bmatrix} \mathbf{P}_{\mathbf{x}\mathbf{x}} & \mathbf{P}_{\mathbf{x}\mathbf{y}} \\ \mathbf{P}_{\mathbf{x}\mathbf{y}}^{\mathrm{T}} & \mathbf{P}_{\mathbf{y}\mathbf{y}} \end{bmatrix}$$
(2)

Broadly speaking, when the prior information is available, typically as high dimensional numerical 120 models in geosciences, see e.g. Carrassi et al. (2018), two main categories of DA (Evensen 2009; 121 Evensen et al. 2022) exists: variational and sequential methods. They both aims at minimizing 122 some energy or functional involving an equation-based dynamical prior and an observation term. 123 Drawing from the link established with Gaussian Processes, we can also consider the case of 124 noisy observations and ease the link with data assimilation formalism, see Särkka and Hartikainen 125 (2012); Särkka et al. (2013); Grigorievskiy et al. (2016). The state space model corresponding to 126 the GP regression problem writes: 127

$$\begin{cases} \mathbf{x}_{t+dt} = \mathbf{M}_{t+dt}\mathbf{x}_t + \boldsymbol{\eta}_t \\ \mathbf{y}_t = \mathbf{H}_t\mathbf{x}_t + \boldsymbol{\varepsilon}_t \end{cases}$$
(3)

where η_t is the m-dimensional noise process and the evolution equation is defined by the feedback linear operator matrix \mathbf{M}_{t+dt} . \mathbf{H}_t is the observation operator at time *t* mapping the state space to the observation space and ε_t the observational error with covariance matrix \mathbf{R}_t . Based on this time-dependent notations, we also consider global observation operator \mathbf{H} with dimensions $(L \times p_k) \times (L \times m)$ and global observational error covariance matrix \mathbf{R} with dimensions $(L \times p_k) \times$ $(L \times p_k)$ as block diagonal matrices whose each block respectively contains the time-dependent observation operator and observational error covariance matrix \mathbf{H}_t and \mathbf{R}_t .

(ii) The Stochastic Partial Differential Equation (SPDE) approach in DA. The Optimal In terpolation implies to factorize dense covariance matrices which is an issue when the size of

spatio-temporal datasets is large. Reduced rank approximations, see e.g. (Cressie and Wikle 2015)
have already been investigated to tackle this specific problem. More recently, the use of sparse
precision matrices has also been proposed by using tapering strategies (Furrer et al. 2006; Bolin
and Wallin 2016) or by making use of the link seen by Lindgren et al. (2011) between Stochastic
Partial Differential Equations (SPDE) and Gaussian Processes. For the latter, if the original link
was made through the Poisson SPDE equation (Whittle 1953):

$$(\kappa^2 - \Delta)^{\alpha/2} X = \tau Z ; \qquad (4)$$

¹⁴³ where $\Delta = \sum_{i=1}^{d} \frac{\partial^2}{\partial s_i^2}$ denotes the Laplacian operator, *Z* is a standard Gaussian white noise, $\kappa = 1/a$, ¹⁴⁴ *a* denotes the range of the GP*X*, $\alpha = \nu + d/2$ and τ relates to the marginal variance of *X*. It can ¹⁴⁵ be extended to more complex linear SPDE involving physical processes like advection or diffusion ¹⁴⁶ (Lindgren et al. 2011; Fuglstad et al. 2015a; Clarotto et al. 2022). The SPDE-based OI formulation ¹⁴⁷ uses precision matrix formalism, as the inverse of the covariance matrix $\mathbf{Q} = \mathbf{\tilde{P}}^{-1} = (dxdy)/\tau^2 \cdot \mathbf{B}^T \mathbf{B}$, ¹⁴⁸ see Eq. (2) where **B** is the discretized version of the fractional differential operator $(\kappa^2 - \Delta)^{\alpha/2}$, ¹⁴⁹ and *dx*, *dy* are the spatial grid step sizes:

$$\mathbf{x}^{\star} = -\mathbf{Q}_{\mathbf{x}\mathbf{x}}^{-1}\mathbf{Q}_{\mathbf{x}\mathbf{y}}\mathbf{y} \tag{5}$$

By construction, **Q** is sparse which means that we solve a system with sparse Cholesky of complexity 150 $O(m^{3/2})$, while the general Cholesky algorithm is of complexity $O(m^3)$. Thus, it opens new avenue 151 to cope with massive observational datasets in geosciences while making use of the underlying 152 physics of such processes. Let note that the so-called SPDE-based approach can also be used 153 as a general spatio-temporal model, even if it is not physically motivated, since it provides a 154 flexible way to handle local anisotropies of a large set of geophysical processes. It has known 155 numerous applications in the past few years, see e.g. Sigrist et al. (2015); Fuglstad et al. (2015b). 156 Though, when considering a spatio-temporal advection-diffusion SPDE, the parameters generally 157 vary continuously across space and/or time making their estimation an additional problem to the 158 original interpolation task. This estimation generally relies on off-line strategies (Fuglstad et al. 159 2015b) embedded in hierarchical models, which can be another computational issue while the set 160 of parameters estimated does not automatically transfer to a new dataset. 161

Regarding the transfer of SPDE formulation in the DA formalism, rewriting the covariance-based
 Eq. (2) in terms of precision matrix leads to:

$$\tilde{\mathbf{P}} = \begin{bmatrix} \mathbf{P}_{\mathbf{x}\mathbf{x}} & \mathbf{P}_{\mathbf{x}\mathbf{x}}\mathbf{H}^{\mathrm{T}} \\ \mathbf{H}\mathbf{P}_{\mathbf{x}\mathbf{x}} & \mathbf{H}\mathbf{P}_{\mathbf{x}\mathbf{x}}\mathbf{H}^{\mathrm{T}} + \mathbf{R} \end{bmatrix}, \quad \tilde{\mathbf{Q}} = \begin{bmatrix} \mathbf{Q}_{\mathbf{x}\mathbf{x}} + \mathbf{H}^{\mathrm{T}}\mathbf{R}^{-1}\mathbf{H} & -\mathbf{H}^{\mathrm{T}}\mathbf{R}^{-1} \\ -\mathbf{R}^{-1}\mathbf{H} & \mathbf{R}^{-1} \end{bmatrix}$$
(6)

¹⁶⁴ which gives an other version of Eq. (5):

$$\mathbf{x}^{\star} = \left(\mathbf{Q}_{\mathbf{x}\mathbf{x}} + \mathbf{H}^{\mathrm{T}}\mathbf{R}^{-1}\mathbf{H}\right)^{-1}\mathbf{H}^{\mathrm{T}}\mathbf{R}^{-1}\mathbf{y}$$
(7)

whose posterior precision matrix of state \mathbf{x}^* is $\mathbf{Q}(\mathbf{x}|\mathbf{y}) = \mathbf{Q}_{\mathbf{xx}} + \mathbf{H}^T \mathbf{R}^{-1} \mathbf{H}$. This type of formulation makes the link between DA framework and SPDE-based formalism maing extensive use of precision matrix formulations.

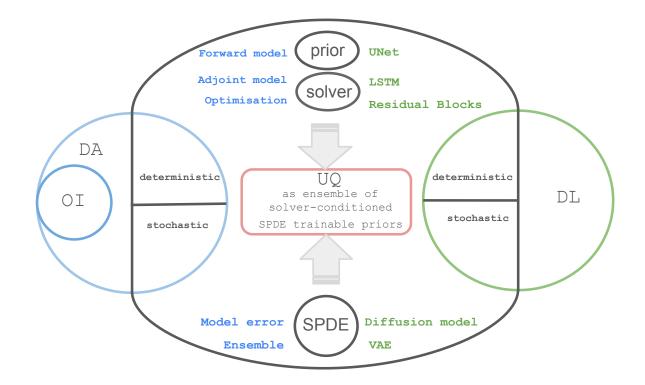


FIG. 1: Combination of SPDE-based formalism, Data Assimilation and Deep Learning to address the challenge of providing UQ in neural variational scheme. We decompose both DA and DL in deterministic and stochastic components to understand how our approach relates to state-of-the-art methods in these fields and how we provide UQ as ensemble of solver-conditioned trainable SPDE priors.

¹⁶⁸ b. Machine and Deep Learning (ML/DL)

From another point of view, deep learning frameworks are currently knowing an intense 169 period of scientific contribution to revisit statistical methods with neural network formulation. 170 The latter enables to use automatic differentiation embedded in the gradient-based optimization 171 as a way of solving traditional inverse problems. Several approaches have been investigated, 172 among others we can think of using DL as a substitute for one component of the DA procedure: 173 surrogate dynamical models making extensive use of state-of-the-art neural networks such as 174 LSTM (Nakamura et al. 2021), UNet (Doury et al. 2023) or Transformers (Tong et al. 2022), 175 reduced order model to compute DA schemes in latent space or with super-resolution component 176 (Brajard et al. 2021), modeling of model/observation error covariance (Cheng and Qiu 2022; 177 Sacco et al. 2022, 2024). The other way rely on end-to-end learning of the entire DA system 178 instead of using DL techniques to address one aspect of DA algorithm among its three main 179 blocks: forward model and its error, observation operator and its error, and DA scheme (EnKF, 180 4DVar, etc.). Such approaches may rely on applying state-of-the-art neural architectures to map 181 observation data to the hidden state sequence, see e.g. UNet (Li et al. 2022), Transformers (Shi 182 et al. 2022) or LSTM architectures (Martin et al. 2023). In particular, when the problem relates 183 to space-time interpolation of partial and noisy observation of geophysical fields, DA-inspired 184 neural schemes have been recently proposed, see e.g. Boudier et al. (2023); Rozet and Louppe 185 (2023); Fablet et al. (2021). The latter specifically suggests a joint learning of prior models and 186 solvers as an end-to-end data-driven variational data assimilation scheme. The so-called 4DVarNet 187 neural scheme is introduced: it involves an implicit iterative gradient-based LSTM solver to 188 minimize a variational cost, close to what is encountered in 4DVar data assimilation (Carrassi 189 et al. 2018). In this variational cost, the dynamical prior is no longer equation-based but is stated 190 as a trainable neural network learnt during the training process. Then, automatic differentiation 191 is used to compute the gradient of the variational cost during the gradient-based iterations, 192 instead of requiring the computation of complex and costly adjoint models (Asch et al. 2016). 193 Drawing from this framework, a neural optimal interpolation scheme has also been proposed 194 (Beauchamp et al. 2022) to reach OI performance with a linear scaling of the solution on the 195 number of space-time variables, leading to a significant speed up in the computation of the solution. 196

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¹⁹⁸ c. Uncertainty Quantification

(i) UQ in DA. Based on its bayesian formalism, Optimal Interpolation provides uncertainty 199 through a posterior covariance matrix $P_{x|y}$. Because the background x^b is generally considered 200 as stationary, this covariance matrix is mainly driven by the sampling of the observation, see 201 e.g. (Zhen et al. 2020), which may not be realistic for dynamical systems. Relying on a similar 202 framework, DA schemes provide the deterministic part of the evolution equation in Eq. 3 as 203 numerical model outputs, and another probabilistic model has to be given for the distribution of 204 the noise. In addition, under Gaussian assumption of the model likelihood, closed forms exist for 205 the mean and standard deviation of the posterior prediction. Ensemble-based sequential schemes 206 also provide a simple way to compute flow-dependant empirical posterior covariance (Song et al. 207 2013) that counteracts the sampling issue used in OI or in most of variational schemes. Hybrid 208 methods in DA (Asch et al. 2016) aims at combining both ensemble-based methods together with 209 variational schemes to benefit from the assets of each methods. 210

(*ii*) UQ for ML approaches. As stated above, DL for UQ generally relies on the estimation of the 211 stochastic components of the DA scheme, namely the covariance model and/or observation errors. 212 More recently, ML schemes aims at substituting the entire DA scheme, either with an explicit 213 model of the posterior distribution, or by a sampling strategy of the latter. On the first approach, 214 because the true posterior is both computationally and analytically intractable, a popular strategy is 215 to estimate an approximate posterior distribution model with trainable parametrization that involves 216 the minimization of the Kullback-Leibler Divergence between the two distributions, which is the 217 similar that maximizing the Evidence Lower Bound (ELBO) (Huang et al. 2019) referred as 218 variational inference (Zhang et al. 2021). More recently, generative models are extensively used 219 to draw samples in the prior distribution. Among them, GANs (Goodfellow et al. 2014), VAEs 220 (Kingma and Welling 2022), normalizing flows (Dinh et al. 2017) and diffusion models (Ho et al. 221 2020) are the most popular. Once a way of sampling the prior is available, the posterior pdf can be 222 obtained after conditioning through a classic DA scheme or method-related conditional generative 223 models. This is the case in score-based data assimilation (Rozet and Louppe 2023) stated as 224 diffusion models where Langevin iterative optimisation scheme can also embed the observation 225 likelihood (Ho et al. 2020) to give a direct access to the posterior. 226

3. Neural variational schemes with SPDE priors

We present here how we draw from the neural variational scheme of Fablet et al. (2021) to embed an augmented state with SPDE prior parametrization. We explain how to parametrize the SPDE with advection-diffusion schemes compliant with many geophysical processes, and how the latter leads to the use of sparse prior precision matrix, which is a key aspect in our use of SPDE priors for surrogate models in DA. A full overview of the neural solver is also presented together with its learning scheme. Last, a presentation of the UQ scheme to generate huge ensembles of the posterior pdf is given.

235 a. Neural MAP solver

Because data assimilation is fundamentally bayesian, most of the methods used to interpolate an observational dataset involve the use of a model prior **x** and the computation of the posterior $E[\mathbf{x}|\mathbf{y}]$ given the observations. For the latter, the computational time might be expensive, even prohibitive, because it implies to solve linear systems with matrices in high dimensions. Recently, alternate solutions have been proposed to compute the posterior. Rather than using linear algebra, we can use a traditional variational data assimilation scheme (Asch et al. 2016) and the state analysis $\mathbf{x}^* = E[\mathbf{x}|\mathbf{y}]$ is obtained by solving the minimization problem:

$$\mathbf{x}^{\star} = \underset{\mathbf{x}}{\operatorname{arg\,min}} \, \mathcal{J}(\mathbf{x})$$

where the variational cost function $\mathcal{J}(\mathbf{x}) = \mathcal{J}_{\Phi}(\mathbf{x}, \mathbf{y}, \Omega)$ is generally the sum of an observation term and a regularization term involving an operator Φ which is typically a dynamical prior:

$$\mathcal{J}_{\Phi}(\mathbf{x}, \mathbf{y}, \Omega) = \mathcal{J}^{o}(\mathbf{x}, \mathbf{y}, \Omega) + \mathcal{J}_{\Phi}^{b}(\mathbf{x})$$
$$= ||\mathcal{H}(\mathbf{x}) - \mathbf{y}||_{\Omega}^{2} + \lambda ||\mathbf{x} - \Phi(\mathbf{x})||^{2}$$
(8)

with \mathcal{H} the observation operator and λ is a predefined or learnable scalar weight. This formulation of functional $\mathcal{J}_{\Phi}(\mathbf{x}, \mathbf{y}, \Omega)$ directly relates to weakly constraint 4DVar, see e.g. Carrassi et al. (2018). When both prior $\Phi(\mathbf{x}) = \mathbf{x}^b$ and observations \mathbf{y} are assumed to be Gaussian with covariance matrices ²⁴⁸ **B** and **R**, $\mathcal{J}_{\Phi}(\mathbf{x}, \mathbf{y}, \Omega)$ can be written:

$$\mathcal{J}_{\Phi}(\mathbf{x}, \mathbf{y}, \Omega) = (\mathbf{H}\mathbf{x} - \mathbf{y})^{\mathrm{T}} \mathbf{R}^{-1} (\mathbf{H}\mathbf{x} - \mathbf{y}) + \lambda \left(\mathbf{x} - \mathbf{x}^{b}\right)^{\mathrm{T}} \mathbf{B}^{-1} \left(\mathbf{x} - \mathbf{x}^{b}\right)$$
(9)

This is well known that equating to zero the gradient of this cost function at a single time t_k produces the exact same analysis formulation that the Kalman Filter analysis step or the BLUE (simple kriging) equations. For inverse problems with time-related processes, the minimization of functional \mathcal{J}_{Φ} usually involves iterative gradient-based algorithms and in particular request to consider the adjoint method in classic equation-based variational data assimilation schemes (Asch et al. 2016) where operator Φ identifies to a deterministic model $\mathbf{x}_{k+1} = \mathcal{M}(\mathbf{x}_k)$:

$$\mathbf{x}^{(i+1)} = \mathbf{x}^{(i)} - \alpha \nabla_{\mathbf{x}} \mathcal{J}_{\Phi}(\mathbf{x}^{(i)}, \mathbf{y}, \Omega)$$

²⁵⁵ Fablet et al. (2020) shows that the so-called 4DVarNet scheme, an end-to-end deep learning ²⁵⁶ framework can be built based on the above variational formulation where both prior operator ²⁵⁷ Φ and Maximum a posterior (MAP) solver Γ , i.e. the operator solving for the gradient-based ²⁵⁸ minimization of the variational cost, are neural networks. For the latter, following meta-learning ²⁵⁹ schemes (Andrychowicz et al. 2016), a residual LSTM-based representation of operator Γ is ²⁶⁰ considered where the *i*th iterative update of the solver is given by:

$$\begin{cases} \mathbf{g}^{(i+1)} = LSTM \left[\alpha \cdot \nabla_{\mathbf{x}} \mathcal{J}_{\Phi}(\mathbf{x}^{(i)}, \mathbf{y}, \Omega), h(i), c(i) \right] \\ \mathbf{x}^{(i+1)} = \mathbf{x}^{(i)} - \mathcal{T} \left(\mathbf{g}^{(i+1)} \right) \end{cases}$$
(10)

with $\mathbf{g}^{(i+1)}$ is the LSTM output using as input gradient $\nabla_{\mathbf{x}} \mathcal{J}_{\Phi}(\mathbf{x}^{(i)}, \mathbf{y}, \Omega)$, while h(i) and c(i)denotes the internal states of the LSTM (Arras et al. 2019), α is a normalization scalar and \mathcal{T} a linear or convolutional mapping.

264

In this formulation, the prior term Φ is jointly estimated so that the reconstruction fulfills at most the loss function \mathcal{L} used during the training process. Typically, such a loss function may be stated 267 as:

$$\mathcal{L}(\mathbf{x}, \mathbf{x}^{\star}) = ||\mathbf{x} - \mathbf{x}^{\star}||^2 + \mathcal{L}_{\text{regul}},$$
(11)

the mean squared error between the reconstruction and the ground truth, with additional regular-268 ization terms, depending on both the specific application and the targeted ouputs of the end-to-end 269 neural scheme. This implies a bi-level optimization scheme in which we refer to Eqs. (8) and (11) 270 resp. as the inner variational cost and the outer training loss function of 4DVarNet schemes. When 271 replacing the outer training loss function by the same inner variational cost, such a scheme would 272 lead to learn the exact 4DVar scheme, but with additional speeding up the optimization process. 273 Instead, using this dual optimization enables to escape from the Gaussian and Linear hypotheses 274 used in the 4DVar formalism so that the training of (11) acts as an additional constraint in the 275 minimization process of (8). 276

277 b. SPDE parametrization

²⁷⁸ Though, the prior Φ is not easily interpretable: it acts as an encoding of the state **x** that helps in ²⁷⁹ the gradient-based minimization process. In this work, we aim at bringing both explainability and ²⁸⁰ stochasticity in the neural scheme by considering as a surrogate model for prior Φ an stochastic ²⁸¹ PDE (SPDE) instead of a given neural-based architecture. The continuous process associated with ²⁸² our state space describes the dynamical evolution of the spatio-temporal prior process $X_{\theta}(\mathbf{s}, t)$ as ²⁸³ a spatio-temporal SPDE embedding the estimation of its parametrization θ , the latter controlling ²⁸⁴ key physical behaviours such as local anisotropy, correlation range, and marginal variance:

$$\mathcal{F}_{t,\boldsymbol{\theta}(t)}\{X(\mathbf{s},t)\} = \tau(\mathbf{s},t)Z(\mathbf{s},t)$$
(12)

where operator $\mathcal{F}_{t,\theta(t)}$ is here considered as linear, then the solution of Eq. (12) is a spatio-temporal Gaussian Process (Lindgren et al. 2011). Regarding the right-hand side noise of the SPDE, it is assumed separable, i.e. $Z(\mathbf{s},t) = Z_t(t) \otimes Z_s(\mathbf{s})$ with $Z_t(t)$ a temporal white noise. $Z_s(\mathbf{s})$ can also be a spatial white noise or we may consider a colored noise to ensure more regularity on process X accross time. Following a state space formalism, the discretization of the stochastic process $X(\mathbf{s}, t)$ is a multivariate gaussian vector \mathbf{x} :

$$\mathbf{x} \sim \mathcal{N}\left(\mathbf{x}^{b}, \mathbf{Q}_{\boldsymbol{\theta}}^{b}\right) \tag{13}$$

where \mathbf{x}^{b} is the deterministic mean of the state, typically *the background*, a coarse approximation of the field in stationary formulations, or the *forecast* in a dynamical data assimilation scheme. \mathbf{Q}^{b}_{θ} is the precision matrix (inverse of the covariance matrix) of the state sequence $\{\mathbf{x}_{0}, \dots, \mathbf{x}_{Ldt}\}$. Because of the explicit link between linear stochastic PDEs and Gaussian Processes, we modify the variational formulation used in 4DVarNet schemes, see e.g. Beauchamp et al. (2023c), by rewriting the matrix formulation of the regularization prior term of Eq. 8 as:

$$\mathbf{x}^{\star} = \operatorname{argmin}_{\mathbf{x}} \mathcal{J}(\mathbf{x}, \mathbf{y}, \Omega) = \operatorname{argmin}_{\mathbf{x}} ||\mathbf{H}\mathbf{x} - \mathbf{y}||_{\Omega}^{2} + \lambda \left[\mathbf{x} - \mathbf{x}^{b}\right]^{\mathrm{T}} \mathbf{Q}_{\theta}^{b} \left[\mathbf{x} - \mathbf{x}^{b}\right]^{\mathrm{T}}$$

²⁹⁹ When $\mathbf{x}^b = \mathbf{0}$ (for simplification) and when denoting **L** the square root of the precision matrix \mathbf{Q}^b_{θ} , ³⁰⁰ the identification with 4DVarNet schemes is direct for $\Phi = (1 - \mathbf{L})$. Indeed, $||\mathbf{x}||^2_{\mathbf{Q}_{\theta}} = \mathbf{x}^T \mathbf{L}^T \mathbf{L} \mathbf{x} =$ ³⁰¹ $||\mathbf{L}\mathbf{x}||^2 = ||\mathbf{x} - \Phi \cdot \mathbf{x}||^2$. In this derived formulation, it is clear that if the SPDE is known, it can ³⁰² be embedded in the inner variational cost used by the LSTM iterative solver to optimize the outer ³⁰³ training loss function. Drawing from the usual neural variational framework, the trainable prior is ³⁰⁴ now SPDE-based, and the parameters θ are embedded in the following augmented state formalism:

$$\tilde{\mathbf{x}} = \begin{bmatrix} \mathbf{x} & \boldsymbol{\theta} \end{bmatrix}^{\mathrm{T}}$$
(14)

The latent parameter θ is potentially non stationary in both space and time and its size is directly related to the size of the data assimilation window *L*.

³⁰⁷ When dealing with geophysical fields, a generic class of non-stationary models generated by ³⁰⁸ stochastic PDEs shall introduce some diffusion and/or advection terms. They are respectively ³⁰⁹ obtained by introducing a local advection operator $\mathbf{m}(\mathbf{s}, t) \cdot \nabla$ where **m** is a velocity field and a local diffusion operator $\nabla \cdot \mathbf{H}(\mathbf{s}, t) \cdot \nabla$ where **H** acts as a two-dimensional diffusion tensor:

$$\frac{\partial X}{\partial t} + \left\{ \kappa^2(\mathbf{s}, t) + \mathbf{m}(\mathbf{s}, t) \cdot \nabla - \nabla \cdot \mathbf{H}(\mathbf{s}, t) \cdot \nabla \right\}^{\alpha/2} X(\mathbf{s}, t) = \boldsymbol{\tau}(\mathbf{s}, t) Z(\mathbf{s}, t)$$
(15)

This way of handling spatio-temporal non-stationarities in SPDE models has been first mentioned in the original paper of Lindgren et al. (2011), but because of the challenge of estimating the full set of space-time parameters, no works have been published to our knowledge pushing this framework into its completeness. Though, some parametrization of purely spatial diffusion process (Fuglstad et al. 2015c) or stationary advection-dominated SPDE (Clarotto et al. 2022) has been successfully applied.

Let stress that the advection-diffusion scheme is a good candidate for many geophysical datasets: this is the case in quasi-geostrophic approximation of Sea Surface Temperature (SST), see e.g. (Ubelmann et al. 2014), but also in the dispersion of atmospheric pollutants (Menut et al. 2021) for instance. Then, this framework provides a generic and convenient way to bring more explainability in terms of space-time covariances of dynamical processes.

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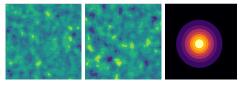
Such a model implies to estimate new parameters $\kappa(\mathbf{s}, t)$, $\mathbf{H}_{2\times2}(\mathbf{s}, t)$ and $\mathbf{m}_{2\times1}(\mathbf{s}, t) = \begin{bmatrix} \mathbf{m}^1 & \mathbf{m}^2 \end{bmatrix}^1$, all varying across space and time along the data assimilation window. In addition, κ needs to be continuous while **m** and **H** additionnaly requires to be continuously differentiable. Regarding the diffusion tensor, we draw from the spatial statistics literature, see e.g. (Fuglstad et al. 2015a), to introduce the scalars $\gamma(\mathbf{s}, t)$, $\beta(\mathbf{s}, t)$, $\mathbf{v}_1(\mathbf{s}, \mathbf{t})$ and $\mathbf{v}_2(\mathbf{s}, \mathbf{t})$ as a generic decomposition of $\mathbf{H}(\mathbf{s}, t)$ through the equation:

$$\mathbf{H}(\mathbf{s},t) = \begin{bmatrix} \mathbf{H}^{1,1} & \mathbf{H}^{1,2} \\ \mathbf{H}^{1,2} & \mathbf{H}^{2,2} \end{bmatrix} (\mathbf{s},t) = \gamma(\mathbf{s},t)\mathbf{I}_2 + \beta(\mathbf{s},t)\mathbf{v}\mathbf{v}^{\mathrm{T}}$$

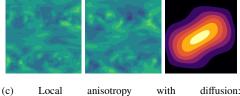
with $\mathbf{v}^{\mathrm{T}} = \begin{bmatrix} \mathbf{v}_{1}(\mathbf{s}) & \mathbf{v}_{2}(\mathbf{s}) \end{bmatrix}$, which models the diffusion tensor as the sum of an isotropic and anisotropic effects, the latter being described by its amplitude and magnitude. This is a valid decomposition for any symmetric positive-definite 2×2 matrix. This leads to the SPDE hyperparametrization $\boldsymbol{\theta}$ of size $m \times L \times 8$ parameters ($\mathbf{H}^{1,2} = \mathbf{H}^{2,1}$): it grows linearly with the potentially ³³³ high dimensional state space. In the end, the SPDE hyperparametrization states as:

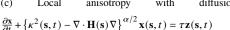
$$\theta = \begin{bmatrix} \kappa & \mathbf{m} & \mathbf{H} & \tau \end{bmatrix}^{\mathrm{T}}$$

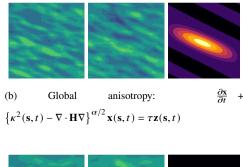
Fig.2 provide some specific examples to show how the modifications in the fractional differential operator leads to more complex spatio-temporal anisotropies. In this four SPDE parametrizations, $\kappa = 0.33$, $\tau = 1$ and $\alpha = 4$.

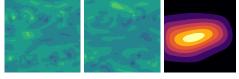


(a) Isotropic model: $\frac{\partial \mathbf{x}}{\partial t} + (\kappa^2(\mathbf{s}, t) - \Delta)^{\alpha/2} \mathbf{x}(\mathbf{s}, t) = \tau \mathbf{z}(\mathbf{s}, t))$









(d) Local anisotropy + Global advection: $\frac{\partial \mathbf{x}}{\partial t}$ + $\{\kappa^2(\mathbf{s}, t) + \mathbf{m} \cdot \nabla - \nabla \cdot \mathbf{H}(\mathbf{s})\nabla\}^{\alpha/2} \mathbf{x}(\mathbf{s}, t) = \tau \mathbf{z}(\mathbf{s}, t)$

FIG. 2: For (a), (b), (c) and (d), one realization of the corresponding SPDE-driven GP at time t = 10 (left panel), t = 20 (middle panel) and covariance (right panel) with central point of domain $\mathcal{D} = [0, 1] \times [0, 1]$

³³⁷ The space-time SPDE is discretized based on a numerical implicit Euler scheme:

$$\frac{\mathbf{x}_{t+dt} - \mathbf{x}_t}{dt} + \mathbf{B}_{t+dt}\mathbf{x}_{t+dt} = \frac{\tau}{\sqrt{dt}}\mathbf{z}_{t+dt}$$

where \mathbf{x}_t and \mathbf{B}_t respectively denote the state space and the finite difference discretization of the fractional differential operator $\frac{\partial}{\partial t} + \{\kappa^2(\mathbf{s},t) + \mathbf{m}(\mathbf{s},t) \cdot \nabla - \nabla \cdot \mathbf{H}(\mathbf{s},t)\nabla\}^{\alpha/2}$ at time $t = 0, \dots, L$. The noise \mathbf{z}_{t+dt} is white in space and $\mathbf{M}_{t+dt} = (\mathbf{I} + dt\mathbf{B}_{t+dt})^{-1}$ denotes the matrix operator that emulates the dynamical evolution of state \mathbf{x} from time t to t + dt. In this case, $\mathbf{T}_{t+dt} = \tau \sqrt{dt} \mathbf{M}_{t+dt}$ corresponds to the dynamical linear model regularized by the product of the noise variance with the square root of the SPDE time step. In case of a more general right-hand term with non-uniform regularization variance { τ_t , t > 0}, and { \mathbf{z}_t , t > 0} are independent realizations of a colored noised driven by a spatial isotropic SPDE:

$$(\kappa_s^2 - \Delta)^{\alpha_s/2} Z_s(\mathbf{s}) = W(\mathbf{s})$$
(16)

with $W(\mathbf{s})$ a white noise with unit variance, the spatial FDM together with time Euler discretization leads to:

$$\mathbf{x}_{t+dt} = \mathbf{M}_{t+dt}\mathbf{x}_t + \mathbf{T}_{t+dt}\mathbf{z}_{t+dt}$$
(17)

where $\widetilde{\mathbf{T}}_{t+dt} = \sqrt{dt} \mathbf{M}_{t+dt} \tau_t \mathbf{L}_s$ and \mathbf{L}_s stands for the Cholesky decomposition of the discretized spatial precision matrix \mathbf{Q}_s introduced of the stochastic process Z_s introduced by Eq. (16). In the case of advection-dominated SPDEs, we involve state-of-the-art upwind schemes (UFDM) for stabilization of the numerical system, by letting the advective transport term, which is the dominating term, collect its information in the flow direction, i.e., upstream or upwind of the point in question. All the calculation details are given in Appendix 5.

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In a compact formulation, using centered finite differences on the diffusion term, and by denoting $\mathbf{a}_{i,j}^{1,t,+} = \max(\mathbf{m}_{i,j}^{1,t}, 0), \mathbf{a}_{i,j}^{1,t,-} = \min(\mathbf{m}_{i,j}^{1,t}, 0), \mathbf{a}_{i,j}^{2,t,+} = \max(\mathbf{m}_{i,j}^{2,t}, 0), \mathbf{a}_{i,j}^{2,t,-} = \min(\mathbf{m}_{i,j}^{2,t}, 0), \text{ the resulting}$ UFDM scheme is:

$$\begin{aligned} \mathbf{x}_{i,j}^{t+1} &= \mathbf{x}_{i,j}^{t} + dt \left[\kappa_{i,j}^{t} \mathbf{x}_{i,j}^{t} + \left(\mathbf{a}_{i,j}^{1,t,+} \mathbf{m}_{i,j}^{1,t,-} + \mathbf{a}_{i,j}^{1,t,-} \mathbf{m}_{i,j}^{1,t,+} \right) + \left(\mathbf{a}_{i,j}^{2,t,+} \mathbf{m}_{i,j}^{2,t,-} + \mathbf{a}_{i,j}^{2,t,-} \mathbf{m}_{i,j}^{2,t,+} \right) \\ &+ \mathbf{H}_{i,j}^{1,1,t} \frac{\mathbf{x}_{i+1,j}^{t} - 2\mathbf{x}_{i,j}^{t} + \mathbf{x}_{i-1,j}^{t}}{dx^{2}} + \mathbf{H}_{i,j}^{2,2,t} \frac{\mathbf{x}_{i,j+1}^{t} - 2\mathbf{x}_{i,j}^{t} + \mathbf{x}_{i,j-1}^{t}}{dy^{2}} \\ &+ \mathbf{H}_{i,j}^{1,2,t} \frac{\mathbf{x}_{i+1,j+1}^{t} - \mathbf{x}_{i+1,j-1}^{t} - \mathbf{x}_{i-1,j+1}^{t} + \mathbf{x}_{i-1,j-1}^{t}}{2dxdy} + \tau_{i,j}^{t} \mathbf{z}_{i,j}^{t+1} \right] \end{aligned}$$

Starting from this numerical scheme, the modified 4DVarNet scheme requires the precision matrix \mathbf{Q}_{θ}^{b} of the state sequence $\{\mathbf{x}_{0}, \dots, \mathbf{x}_{Ldt}\}$. Here, $\mathbf{x}_{0} \sim \mathcal{N}(\mathbf{0}, \mathbf{P}_{0})$ denotes the initial state and $\mathbf{Q}_{0} = \mathbf{P}_{0}^{-1}$ is always taken as the precision matrix obtained after a given stabilization run, i.e. the evolution of the dynamical system over N timesteps using as stationary parameters the initial parametrization θ_0 of the SPDE at time t = 0, then we can rewrite :

$$\{\mathbf{x}_0,\cdots,\mathbf{x}_{Ldt}\}=\mathbf{M}_G\left[\begin{array}{c}\mathbf{x}_0\\\mathbf{z}\end{array}\right]$$

with $\mathbf{z} = [\mathbf{z}_1, \dots, \mathbf{z}_t]^{\mathrm{T}}$ and

$$\mathbf{M}_{G} = \begin{bmatrix} \mathbf{I} & 0 & 0 & 0 & 0 & \cdots & 0 \\ \mathbf{M}_{1} & \mathbf{T}_{1} & 0 & 0 & 0 & \cdots & 0 \\ \mathbf{M}_{2}\mathbf{M}_{1} & \mathbf{M}_{2}\mathbf{T}_{1} & \mathbf{T}_{2} & 0 & 0 & \cdots & 0 \\ \mathbf{M}_{3}\mathbf{M}_{2}\mathbf{M}_{1} & \mathbf{M}_{3}\mathbf{M}_{2}\mathbf{T}_{1} & \mathbf{M}_{3}\mathbf{T}_{2} & \mathbf{T}_{3} & 0 & \cdots & 0 \\ \vdots & \ddots & \ddots & \ddots & \ddots & \ddots & 0 \\ \vdots & \ddots & \ddots & \ddots & \ddots & \ddots & \ddots & 0 \\ \vdots & \ddots & \ddots & \ddots & \ddots & \ddots & \mathbf{T}_{L} \end{bmatrix}$$

With the additional notation $\mathbf{S}_k = \mathbf{T}_k \mathbf{T}_k^{\mathrm{T}}$, see Eq. (17), the precision matrix \mathbf{Q}^b writes, see Appendix c for all the details:

$$\mathbf{Q}_{\boldsymbol{\theta}}^{b} = \frac{1}{dt} \begin{bmatrix} \mathbf{P}_{0}^{-1} + \widetilde{\mathbf{Q}}_{s,1} & -\widetilde{\mathbf{Q}}_{s,1} \mathbf{M}_{1}^{-1} & 0 & 0 & 0 & \dots & 0 \\ -(\mathbf{M}_{1}^{\mathrm{T}})^{-1} \widetilde{\mathbf{Q}}_{s,1} \mathbf{M}_{1}^{\mathrm{T}} \widetilde{\mathbf{Q}}_{s,1} \mathbf{M}_{1} + \widetilde{\mathbf{Q}}_{s,2} & -\widetilde{\mathbf{Q}}_{s,2} \mathbf{M}_{2}^{-1} & 0 & 0 & \dots & 0 \\ 0 & -(\mathbf{M}_{2}^{\mathrm{T}})^{-1} \widetilde{\mathbf{Q}}_{s,2} & \mathbf{M}_{2}^{\mathrm{T}} \widetilde{\mathbf{Q}}_{s,2} \mathbf{M}_{2} + \widetilde{\mathbf{Q}}_{s,3} & -\widetilde{\mathbf{Q}}_{s,3} \mathbf{M}_{3}^{-1} & 0 & \dots & 0 \\ 0 & -(\mathbf{M}_{2}^{\mathrm{T}})^{-1} \widetilde{\mathbf{Q}}_{s,2} & \mathbf{M}_{2}^{\mathrm{T}} \widetilde{\mathbf{Q}}_{s,2} \mathbf{M}_{2} + \widetilde{\mathbf{Q}}_{s,3} & -\widetilde{\mathbf{Q}}_{s,3} \mathbf{M}_{3}^{-1} & 0 & \dots & 0 \\ 0 & -(\mathbf{M}_{2}^{\mathrm{T}})^{-1} \widetilde{\mathbf{Q}}_{s,2} & \mathbf{M}_{2}^{\mathrm{T}} \widetilde{\mathbf{Q}}_{s,2} \mathbf{M}_{2} + \widetilde{\mathbf{Q}}_{s,3} & -\widetilde{\mathbf{Q}}_{s,3} \mathbf{M}_{3}^{-1} & 0 & \dots & 0 \\ 0 & \ddots & \ddots & \ddots & \ddots & \ddots & \ddots & 0 \\ 0 & \ddots & \ddots & \ddots & \ddots & \ddots & \ddots & 0 \\ 0 & \ddots & \ddots & \ddots & \ddots & \ddots & \ddots & 0 \\ 0 & \ddots & \ddots & \ddots & \ddots & 0 & -(\mathbf{M}_{L-1}^{\mathrm{T}})^{-1} \widetilde{\mathbf{Q}}_{s,L-1} \mathbf{M}_{L-1}^{\mathrm{T}} + \widetilde{\mathbf{Q}}_{s,L} - \widetilde{\mathbf{Q}}_{s,L} \mathbf{M}_{L}^{-1} \\ 0 & \ddots & \ddots & \ddots & 0 & -(\mathbf{M}_{L}^{\mathrm{T}})^{-1} \widetilde{\mathbf{Q}}_{s,L} & \mathbf{M}_{L}^{\mathrm{T}} \widetilde{\mathbf{Q}}_{s,L} \mathbf{M}_{L} \end{bmatrix}$$

$$(18)$$

where $\widetilde{\mathbf{Q}}_{s,t}$ is the precision matrix of the colored noise weighted by the non-uniform regularization variance τ_t . As clearly visible, the sparsity of \mathbf{Q}_{θ}^b is high, which is key in traditional SPDE-based GP inference, but also in our approach, see Section c.

c. Neural solver with augmented state

³⁷⁰ Overall, let denote by $\Psi_{\theta,\Gamma}(\tilde{\mathbf{x}}^{(0)},\mathbf{y},\Omega)$ the output of the end-to-end learning scheme given the ³⁷¹ SPDE-based dynamical model with parameters θ and the neural residual architecture for the solver ³⁷² Γ , see Fig. 3 and Algorithm 1, the initialization $\tilde{\mathbf{x}}^{(0)}$ of augmented state $\tilde{\mathbf{x}}$ and the observations \mathbf{y} ³⁷³ on domain Ω .

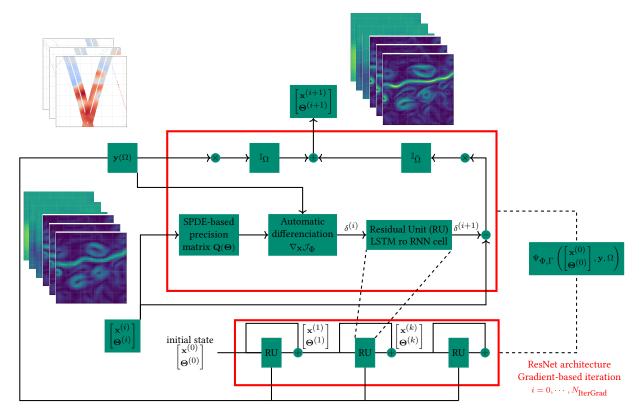


Fig. 3: Sketch of the gradient-based algorithm. I_{Ω} acts as a masking operator for any spatiotemporal location not in Ω .

Algorithm 1: Variational scheme with SPDE-based GP prior and implicit neural solver

Data:

 $\mathbf{x} \in \mathbb{R}^{T \times m} = {\mathbf{x}_k}, \ k = 1, \cdots, T$

 $\mathbf{y}_{\Omega} = {\mathbf{y}_{k,\Omega_k}}, \ k = 1, \cdots, T$: observations on domains $\Omega_k \subset \mathcal{D}$

 N_I : number of iterations

 η : gradient step

Init:

 $\tilde{\mathbf{x}}^{\star,(0)}$

List of procedures:

Train_ $\Psi_{\theta,\Gamma}$: end-to-end learning procedure with:

 θ : parameter of the SPDE-based prior operator;

GradLSTM: residual NN-based representation of $\nabla_{\mathbf{x}} \mathcal{J}(\mathbf{x})$

 Γ : iterative gradient-based update operator:

i = 0

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while $i < N_I$ do

$$\mathbf{Q}^{b}_{\boldsymbol{\theta}^{\star,(i)}} = \mathbf{P}^{-1}_{G}(\boldsymbol{\theta}^{\star,(i)}) = \mathbf{M}^{-1}_{G}(\boldsymbol{\theta}^{\star,(i)})^{\mathrm{T}} \begin{bmatrix} \mathbf{P}^{-1}_{0} & 0 & \dots & 0 \\ 0 & \mathbf{I} & \dots & 0 \\ \vdots & \ddots & \ddots & \vdots \\ 0 & 0 & \dots & \mathbf{I} \end{bmatrix} \mathbf{M}^{-1}_{G}(\boldsymbol{\theta}^{\star,(i)})$$

$$\tilde{\mathbf{x}}^{(i+1)} \leftarrow \tilde{\mathbf{x}}^{(i)} - \eta \times GradLSTM(\tilde{\mathbf{x}}^{(i)})$$

$$N_{I} \nearrow ; \eta \searrow ; i \leftarrow i+1$$

end

end

Result: $\tilde{\mathbf{x}}^{\star} \leftarrow \Psi_{\boldsymbol{\theta},\Gamma}(\tilde{\mathbf{x}}^{(0)}, \mathbf{y}, \Omega)$

Then, the joint learning for the weights $\omega_{\theta,\Gamma}$ of the neural scheme given the SPDE formulation that is chosen (isotropic or not, non-stationary or not, etc.) and the architecture of operator Γ is stated as the minimization of the mixed loss function $\mathcal{L}(\mathbf{x}, \mathbf{x}^{\star}, \boldsymbol{\theta}^{\star})$, fully explained in Section d:

$$\omega_{\Psi_{\theta,\Gamma}}^{\star} = \arg\min_{\theta,\Gamma} \left[\mathcal{L}(\mathbf{x}, \mathbf{x}^{\star}, \theta^{\star}) \right] \text{ s.t. } \tilde{\mathbf{x}}^{\star} = \Psi_{\theta,\Gamma}(\tilde{\mathbf{x}}^{(0)}, \mathbf{y}, \Omega)$$
(19)

(*i*) Initialization of the augmented state. The initial state $\tilde{\mathbf{x}}^{(0)}$ should be filled with 0 in the nonobserved part of domain \mathcal{D} . Because there is obviously no observation of the SPDE parameters, we might guide the training process at the first iteration of the solver. This proved to be particularly helpful when dealing with realistic geophysical datasets for which the advection diffusion scheme is meaningful. In that case, see for instance Application 2 on real SSH datasets in Section 4, we used first and second-order derivatives of the initial state $\mathbf{x}^{(0)}$ as initial parametrizations for the advection and diffusion process:

$$\mathbf{m}^{(0)} = \begin{pmatrix} \frac{\partial \mathbf{x}^{(0)}}{\partial x} \\ \frac{\partial \mathbf{x}^{(0)}}{\partial y} \end{pmatrix}, \text{ and } \mathbf{v}^{(0)} = \begin{pmatrix} \frac{\partial^2 \mathbf{x}^{(0)}}{\partial x^2} \\ \frac{\partial^2 \mathbf{x}^{(0)}}{\partial y^2} \end{pmatrix}$$

 $\kappa^{(0)}$ and $\tau^{(0)}$ are both using absolute values of the normalized gradient norms, while $\beta^{(0)}$ and $\gamma^{(0)}$ are resp. set to 1 and 0. Because κ , τ and γ are both strictly positive, see Section b, we use ReLu activation function on these three parameters to ensure their consistency.

(*ii*) Computational aspects. The sparse formulation of the precision matrix \mathbf{Q}_{θ}^{b} is key in the memory-saving component of the algorithm because the latter relies on a set of N_{I} gradient-based iterations, meaning that for a single interpolation task, the precision matrix \mathbf{Q}_{θ}^{b} is stored N_{I} times along the computational graph with updated values of the SPDE parameters $\boldsymbol{\theta}^{(i)}$.

392 d. Learning scheme

³⁹³ (*i*) *Training loss*. The joint problem of estimating the best reconstruction and infering realistic ³⁹⁴ SPDE parametrizations is difficult because according to the size and nature of the dataset, the ³⁹⁵ spatio-temporal interpolation may not always benefit from knowing the exact set of true SPDE ³⁹⁶ parameters. Indeed, if the degree of sparsity of the observation dataset is low, the reconstruction ³⁹⁷ may be good despite a poor estimation of the covariance matrix. In the other way, if the degree ³⁹⁸ of sparsity is high, much more difficult will be the estimation of the SPDE underlying parameters. In this section, we benefit from the supervised configuration of neural variational scheme to train this joint problem. In this Section, we show how to embed the global precision matrix defined by Eq. 18 in our neural scheme and define which training loss is the more appropriate to handle the bi-level optimization scheme (of both the inner variational cost and the training loss).

⁴⁰³ For the training process, we may consider different loss functions:

•
$$\mathcal{L}_1(\mathbf{x}, \mathbf{x}^*) = ||\mathbf{x} - \mathbf{x}^*||^2$$
 is the L2-norm of the difference between state \mathbf{x} and reconstruction \mathbf{x}^*

• $\mathcal{L}_2(\mathbf{x}, \boldsymbol{\theta}^{\star}) = -|\mathbf{Q}_{\boldsymbol{\theta}^{\star}}^b| + \mathbf{x}^{\mathrm{T}} \mathbf{Q}_{\boldsymbol{\theta}^{\star}}^b \mathbf{x}$ is the negative log-likelihood of the true states given the estimated precision matrix $\mathbf{Q}_{\boldsymbol{\theta}^{\star}}^b$, thus ensuring consistency between the actual ground truth and the SPDE parameters.

Using \mathcal{L}_1 will lead to satisfactory reconstructions without any constraints on the SPDE parameters. The single use of \mathcal{L}_2 should lead to satisfactory results if the analytical solver, i.e. the inversion of the linear system or the gradient-based minimization of the variational cost, were used. But because the solver is trained, it also needs to be constrained by an appropriate loss function for the reconstruction, meaning \mathcal{L}_1 . The best solution is then to create a mixed loss function, combination of \mathcal{L}_2 to estimate at best the SPDE parameters and optimize the prior model, and \mathcal{L}_1 to satisfy the reconstruction criteria and optimize the solver.

The log-determinant of the precision matrix $\log |\mathbf{Q}_{\theta^{\star}}^{b}|$ is usually difficult to handle when computing \mathcal{L}_{2} . Hopefully, based on the particular structure and the notations already introduced for the spatio-temporal precision matrix \mathbf{Q}_{θ}^{b} in Eq.(B2), it writes, see e.g. Clarotto et al. (2022):

$$\log |\mathbf{Q}_{\theta^{\star}}^{b}| = \log |\mathbf{P}_{0}^{-1}| + \log |\mathbf{S}_{1}^{-1}| + \dots + \log |\mathbf{S}_{L}^{-1}|$$

$$= \log |\mathbf{P}_{0}^{-1}| + \sum_{i=1}^{L} \log \left(|\mathbf{L}_{i}\mathbf{L}_{i}^{\mathrm{T}}| \right)$$

$$= \log |\mathbf{P}_{0}^{-1}| + 2\sum_{i=1}^{L} \sum_{j=1}^{m} \log \mathbf{L}_{i}(j, j)$$
(20)

where L_i denotes here the Cholesky decomposition of S_k^{-1} . In case of unsupervised learning, the same strategy may apply but \mathcal{L}_2 will be the likelihood of the observations given the estimated SPDE parameters since in this case, the true states would not be available during the training process.

(*ii*) Two-step learning schemes. In the initial version of 4DVarNet schemes, the output \mathbf{x}^* provided 421 by the neural formulation is deterministic and can be seen as the posterior mean of the state given 422 the observations. Within this SPDE-based parametrization of the prior \mathbf{x}^b , we also aim at providing 423 the distribution of the prior as a GP process. Though, for realistic geophysical fields, even the prior 424 cannot be considered as a zero-mean Gaussian field and some non-linearities have to be accounted 425 for in its deterministic mean \mathbf{x}^b . To solve for this specific case, we involve a 2-step learning process 426 in which the deterministic mean \mathbf{x}^b is first estimated by a 4DVarNet scheme applied on coarser 427 resolution than the actual observations. Second, the modified 4DVarNet scheme is involved to 428 estimate jointly the posterior mean \mathbf{x}^{\star} together with the SPDE parametrization θ^{\star} of the prior 429 $\mathbf{x} \sim \mathcal{N}(\mathbf{x}^b, \mathbf{Q}^b_{\theta^{\star}})$. This two-steps procedure also enables to simplify the SPDE training scheme: the 430 first guess is estimated based on the entire set of observations along the state sequence, while the 431 SPDE parametrization is estimated only on a reduced window of length 5, centered on the targeted 432 time of interest. This considerably reduces the size of precision matrix \mathbf{Q}^b_{θ} , making the algorithm 433 scalable for any application. Fig. 4 shows a schematic overview of this two-steps learning scheme 434 with illustrations coming from realistic Sea Surface Heigh datasets provided in Application 2, 435 Section 4. In the end, this will lead to a stochastic version of the neural variational scheme based 436 on the prior GP distribution in which we can sample ensemble members. 437

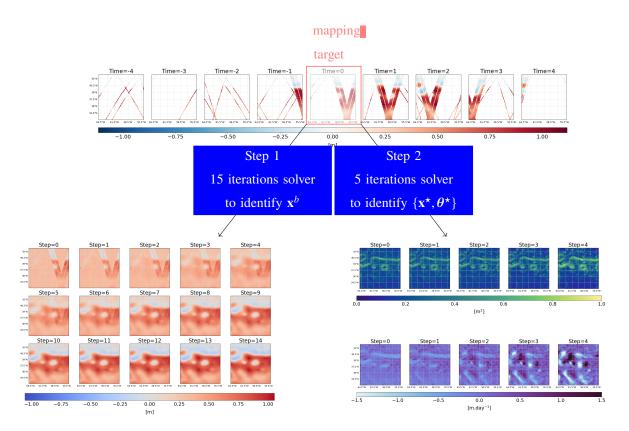


FIG. 4: Two-step adaptation of the 4DVarNet scheme: on the top panel, an example of 4 nadirs+SWOT observations along the 9-day data assimilation window, see Application 2 in Section 4. The target reconstruction day is at the center of the DAW. On the bottom-left panel is displayed the 15 iterations of the first 4DVarNet solver to retrieve the prior mean \mathbf{x}^b : at the beginning, the non-observed parts of the domain are filled with the global mean of the training dataset (0 when normalized). On the bottom-right panel is displayed the 5 iterations of the modified 4DVarNet solver with augmented state formalism to retrieve both \mathbf{x}^* and θ^* : at the beginning, the parameters in θ are initialized with partial derivatives of \mathbf{x}^b , see Section 3.c

(iii) Complementary PyTorch developments. Regarding the implementation of our model, we 438 use Pytorch (Paszke et al. 2017) whose sparse linear algebra is not providing yet a sparse Cholesky 439 algorithm and a sparse solver of linear systems, which is critical, especially when computing the 440 likelihood \mathcal{L}_2 in the training loss function. As a consequence, despite the theoretical tools have 441 been fully detailed in the previous sections, we had to implement new functionalities based on 442 scipy sparse linear algebra (Virtanen et al. 2020) to store the precision matrices in an efficient 443 way and compute the inner variational cost, see again Eq.8. In particular, we provide a PyTorch 444 extension of sparse cholesky matrices, see Appendix c, based on the CHOLMOD supernodal 445 Cholesky factorization (Chen et al. 2008) and draw from Seeger et al. (2019) to provide the 446

⁴⁴⁷ backward pass of the sparse Cholesky decomposition.

449 e. UQ scheme

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Because the PDE is stochastic, it provides an easy way to generate a set of *N* gaussian prior simulations \mathbf{x}_i , $i = 1, \dots, N$:

$$\mathbf{x}_i = \mathbf{x}^b + \mathbf{L}_{\boldsymbol{\theta}}^b \mathbf{z}_i$$

where \mathbf{L}_{θ}^{b} stands for the Cholesky decomposition of \mathbf{Q}_{θ}^{b} and \mathbf{z}_{i} is a white noise. From a geostatistical 452 point of view, this can be seen as SPDE-based spatio-temporal non conditional simulations of state 453 \mathbf{x} , meaning that we produce surrogate simulations sharing the same physical properties than the true 454 states. These simulations are then conditioned by the neural solver given the observations available 455 y, see Fig. 5. To do so, we draw from traditional geostatistics to realize SPDE-based spatio-456 temporal non conditional simulations with a kriging-based conditioning (Wackernagel 2003). 457 Except that we replace the kriging algorithm by our neural approach, which has to be seen as a 458 generic interpolation tool here: 459

$$\mathbf{x}_{i}^{\star}(\mathbf{s},t) = \mathbf{x}^{\star}(\mathbf{s},t) + (\mathbf{x}_{i}(\mathbf{s},t) - \widehat{\mathbf{x}}_{i}(\mathbf{s},t))$$
(21)

where \mathbf{x}^{\star} denotes the neural-based interpolation, \mathbf{x}_i is one SPDE non-conditional simulation of the process \mathbf{x} based on the parameters θ^{\star} and $\mathbf{\hat{x}}_i$ is the neural reconstruction of this non-conditional simulation, using as pseudo-observations a subsampling of \mathbf{x}_i based on the actual data locations. Because $\mathbf{E}[\mathbf{x}_i - \mathbf{\hat{x}}_i] = 0$, the resulting simulation is well conditioned by the observations at data locations.

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⁴⁶⁶ Running an ensemble of *N* conditional simulations gives an approximation of the probability ⁴⁶⁷ distribution function $p_{\mathbf{x}|\mathbf{y}}$ of state $\mathbf{x}^{\star} = \mathbf{x}|\mathbf{y} = {\mathbf{x}_0|\mathbf{y}, \dots, \mathbf{x}_L|\mathbf{y}}$. The ensemble mean $\overline{\mathbf{x}_i^{\star}}$ will be \mathbf{x}^{\star} in the limits of $N \to +\infty$:

$$\frac{1}{N} \sum_{i} \mathbf{x}_{i}^{\star}(\mathbf{s}, t) \xrightarrow[N \to +\infty]{} \mathbf{x}^{\star}(\mathbf{s}, t)$$

Such an approach has already been successfully tested in Beauchamp et al. (2023a) when using analog operator strategy (Tandeo et al. 2015) to draw non-conditional simulation in the prior distribution.

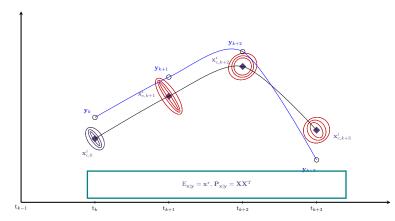


FIG. 5: Ensemble-based neural variational scheme with SPDE-based GP prior: the assimilation is not sequential. The inversion scheme embeds the global precision matrix of the state sequence $\{\mathbf{x}_0, \dots, \mathbf{x}_{Ldt}\}$ in the inner variational cost to minimize. Ensemble members are generated from the Gaussian prior surrogate SPDE model, then conditioned by the neural solver so that the posterior pdf is no longer Gaussian. The posterior pdf is empirically ensemble-based computed: $\mathbf{P}_{\mathbf{x}|\mathbf{y}} = \mathbf{X}\mathbf{X}^{\mathrm{T}}$ with $\mathbf{X} = (1/\sqrt{N-1}) \left[\mathbf{x}_i^{\star} - \mathbf{x}^{\star} \right]$

Let note that in this formulation, the idea is to run SPDE-based conditional simulation of the prior 472 Matérn field $X(\mathbf{s},t)$. One ensemble member is obtained by running one space-time simulation, 473 and two space-time neural reconstructions along the data assimilation window, then combined 474 through Eq. (21). As a consequence, there is no sequential assimilation. Though, in the idea, 475 such a conditioning is exactly similar to the one produced by EnKF simulations: in geostatistical 476 terms we can interpret the forecast step of the EnKF as being unconditional simulations at 477 time t-1 generating N realizations of a non-stationary random function (SPDE-based here) 478 for time t. Both prior mean and covariance matrix are then computed directly on this set of 479 unconditional realizations, before the analysis step, i.e. their conditiong with the observations. 480 The key advantage of the EnKF is its flow-dependency that propagates the uncertainties at each 481

time step with the evolution model, while the classic EnOI method generally used by geoscientists, 482 see e.g. Asch et al. (2016); Counillon and Bertino (2009) replaces the flow-dependent EnKF 483 error covariance matrix by a stationary matrix calculated from an historical ensemble. This is an 484 important difference with our approach: while the neural architecture can be seen as a way to learn 485 neural ensemble-based optimal interpolation models and solvers, our GP prior encoded by its 486 SPDE precision matrix, built according to the FDM scheme with varying parameters over space 487 and time, still allows for flow dependency, based on the learning of the SPDE parametrization 488 given the input observations. 489

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Such a strategy also shares many similarities with the so-called conditional generative models 491 in deep learning, see (Goodfellow et al. 2014; Kingma and Welling 2022; Dinh et al. 2017; Ho 492 et al. 2020). Instead of using neural networks to learn how to simulate in the prior distribution, 493 we use here a generic class of advection-diffusion SPDE which we see as a first step towards 494 physically-sounded generatives models in learning-based methods. Fig. 6 displays an example 495 of such non-conditional simulations for the term $\mathbf{x} - \mathbf{x}^b$, again for Application 2 on realistic SSH 496 datasets over some energetic area along the Gulf Stream, see Section 4, i.e. the anomaly between 497 the true state sequence and the deterministic mean \mathbf{x}^b , retrieved from a preliminary 4DVarNet 498 coarse resolution scheme. We can appreciate how both the generic class of SPDE selected here, 499 together with the training of its parameter, leads to realistic anomalies with a clear increase of 500 the variance along the main meander of the Gulf Stream, due to a correct distribution of SPDE 501 parameter $\tau^{\star}(\mathbf{s}, t)$. For a detailed analysis of the parameter estimation, please report to Section 4. 502

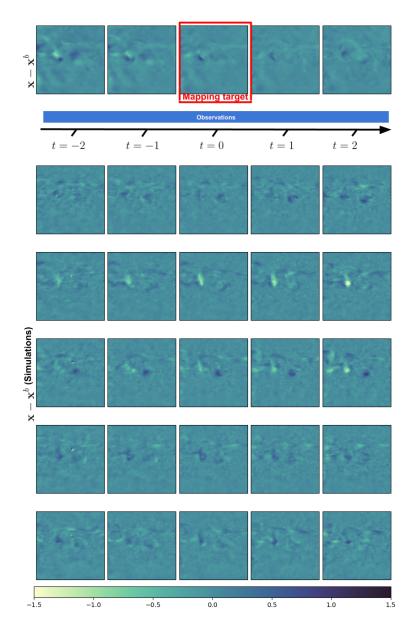


FIG. 6: Top panel: an example of the Ground Truth anomaly between $\mathbf{x} - \mathbf{x}^b$; Bottom panel: five simulations of the same anomaly based on the SPDE parametrization θ learned after training

503 **4. Results**

⁵⁰⁴ In this Section, we provide two applications of this work:

• The first example relies on a spatio-temporal GP simulation driven by a non-stationary spatial diffusion tensor. Because the dynamical process is linear, the best reconstruction is provided

⁵⁰⁷ by the optimal interpolation using the SPDE parameters used in the simulation to fill in the ⁵⁰⁸ precision matrix.

The second example uses an Observation System Simulation Experiment (OSSE) of the Sea Surface Height (SSH) along the Gulf Stream. We will use the SPDE-based prior as a surrogate model along the data assimilation window to provide ensemble members of the posterior distribution.

⁵¹³ a. Diffusion-based non stationary GP

Dataset. In this first application, we simulate 500 states of a GP driven by the following diffusion
 SPDE:

$$\left\{\kappa^{2} - \nabla \cdot \mathbf{H}(\mathbf{s})\nabla\right\}^{\alpha/2} \mathbf{x}(\mathbf{s},t) = \tau \mathbf{z}(\mathbf{s},t)$$
(22)

The regularity parameter $\kappa = 0.33$ is fixed over space and time. To ensure the GP to be smooth enough, we use a value of $\alpha = 4$. Such a formulation enables to generate GPs driven by local anisotropies in space leading to non stationary spatio-temporal fields with eddy patterns. The diffusion tensor **H** is a 2-dimensional diffusion tensor generated by drawing from the spatial statistics literature, see e.g. (Fuglstad et al. 2015a). We introduce a generic decomposition of **H**(**s**, *t*) through the equation:

$$\mathbf{H} = \gamma \mathbf{I}_2 + \beta \mathbf{v}(\mathbf{s})^{\mathrm{T}} \mathbf{v}(\mathbf{s})$$

with $\gamma = 1$, $\beta = 25$ and $\mathbf{v}(\mathbf{s}) = (v_1(\mathbf{s}), v_2(\mathbf{s}))^T$ using a periodic formulation of its two vector fields components, see Section b. We use the Finite Difference Method in space coupled with an implicit Euler scheme in time to solve the equation. Let $\mathcal{D} = [0, 100] \times [0, 100]$ be the square spatial domain of simulation and $\mathcal{T} = [0, 500]$ the temporal domain. Both spatial and temporal domains are discretized so that the simulation is made on a uniform Cartesian grid consisting of points (x_i , y_j , t_k) where $x_i = i\Delta x$, $y_j = j\Delta j$, $t_k = k\Delta t$ with Δx , Δy and Δt all set to 1.

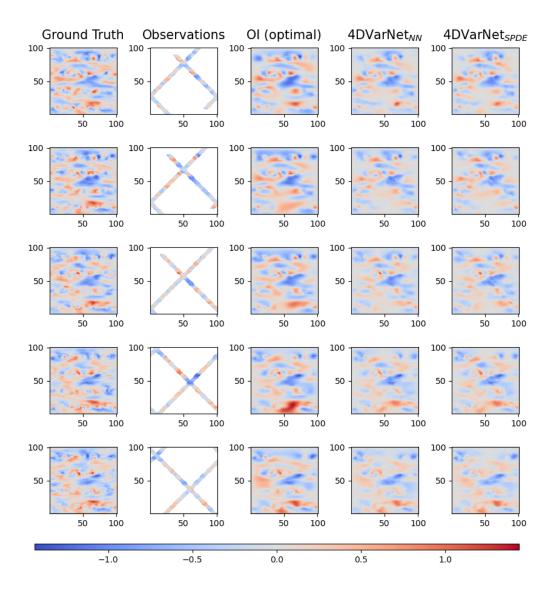


FIG. 7: From left to right: Ground Truth, Pseudo-observations, Optimal Interpolation, neural variational scheme with UNet-based and SPDE-based prior operator Φ . A data assimilation window of length 5 is used

Training setting. We train the neural architectures for both UNet-based and SPDE-based priors with Adam optimizer on 50 epochs. The training period goes from timestep 100 to 400. During the training procedure, we select the best model according to metrics computed over the validation period from timestep 30 to 80. Overall, the set of metrics is computed on a test period going from timestep 400 to 500. No further improvements in the training losses are seen when training the model longer. We use a data assimilation window of length 5 and generate pseudo-observations
 from the ground truth, inspired by orbiting satellites tracks around the earth, see Application 2 in
 Section b.

536

Results. Fig. 7 displays the results obtained by Optimal Interpolation and the neural implicit 537 solver with both UNet and SPDE-based parametrization of the prior. Neural-based reconstructions 538 are optimized at the center of the assimilation window (T = 2), which is why the performance 539 may be affected for other leadtimes. There is no significant differences between the two prior 540 formulations used in the neural scheme, which overall retrieve the main patterns of the OI. Some 541 artefacts may appear due to the observation term in the inner variational cost that tends to lead the 542 solution towards the observation in its close neighbourhood. Improvements may be expected when 543 adding regularization terms in the training loss to counteract such effects, see e.g. Beauchamp 544 et al. (2023b). Regarding the derived framework proposed in this work, one of the question was: 545 is it possible to retrieve interpretable SPDE parametrizations from the joint learning setting? Two 546 configurations were considered: when using as initial condition for the parametrization a gradient-547 based information of the accumulated alongtrack observations, i.e. $\mathbf{H}_{11} = \nabla_{\vec{x}} \mathbf{y}$ and $\mathbf{H}_{22} = \nabla_{\vec{y}} \mathbf{y}$; or 548 an isotropic initialization, i.e. $\mathbf{H} = \mathbf{I}$. The first configuration enables to identify patterns in zonal 549 and meridional components of the true diffusion tensor and leads to an optimal parametrization 550 θ^{\star} very close to the true diffusion tensor. While leading to different SPDE parametrizations θ^{\star} , 551 the interpolation metrics are similar in the end for the two initializations. In addition, using an 552 isotropic initial condition is more general (see the next realistic SSH application) and also retrieves 553 in the end the main zonal flow directions encoded by H_{22} , while the meriodional and periodic 554 structures of \mathbf{H}_{11} and \mathbf{H}_{12} are partly seen as well. 555

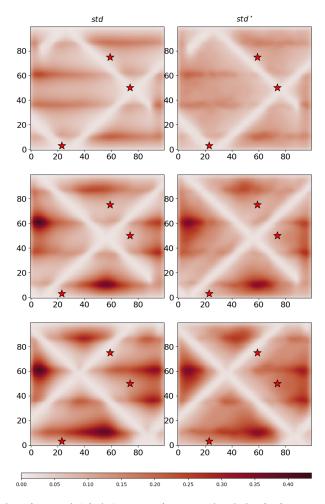


FIG. 8: True (left) and estimated (right) posterior standard deviations at the beginning, center and end of the assimilation window time = 0, 2, 4

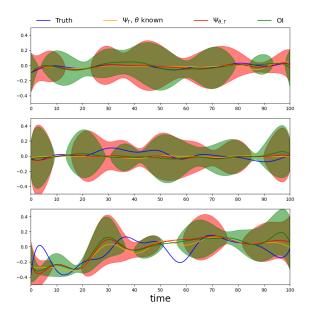


FIG. 9: Ground truth, OI and its posterior variance (blue), neural scheme with SPDE parametrization θ known (orange) and with inference of θ (red) for the same three points identified in Fig. 8 along the 100 time step test period

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Because we know the process is Gaussian, we know that $\mathbf{Q}(\mathbf{x}|\mathbf{y}) = \mathbf{Q}^b + \mathbf{H}^T \mathbf{R}^{-1} \mathbf{H}$, see Section 558 2, to compute the closed form of the posterior pdf. Looking at the estimated posterior variance 559 obtained starting from isotropic initial condition of the parameters, see Fig. 8, we understand that 560 even if the initial set of parameters is not retrieved, they still remain interpretable. Visually, a sim-561 ulation produced by this set of SPDE parameters is also consistent and propose a spatio-temporal 562 diffusion process close to the original one. It also highlights the dependency of the OI uncertainty 563 quantification to the sampling scheme, especially here where no noise is added to the partial set 564 of observations. On Fig. 9 is also shown the time series on three different locations (red stars) 565 $\in \mathcal{D}$, see again Fig. 8, of the GT (blue line) along the test period (100 time steps), the OI and 566 corresponding standard deviation (green line), and the neural scheme estimations when the SPDE 567 parametrization is known (orange line) or estimated (red line with associated uncertainty). Both 568 neural scheme configurations are generally close which validates the capability of this framework 569 to estimate jointly both state and prior parametrization. Because the neural scheme is optimized 570 on the global MSE, its solution may deviate more or less significantly from the OI depending of 571 the position we are looking at in domain \mathcal{D} . 572

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574

Last, Fig. 10 provides the scatterplot of the global MSE w.r.t the OI variational cost:

$$\mathcal{J}_{OI}(\mathbf{y}, \mathbf{x}, \boldsymbol{\theta}^{\star}) = ||\mathbf{y} - \mathbf{H}\mathbf{x}||^2 + \mathbf{x}^{\mathrm{T}} \mathbf{Q}_{\boldsymbol{\theta}^{\star}}^b \mathbf{x}$$

throughout the iteration process after training of the neural schemes. For SPDE-based prior, the 575 initial parametrization $\theta^{(0)}$ relates to an isotropic GP process. LSTM-based iterative solvers are all 576 consistent with the optimal solution in terms of MSE. When the latter is used as training loss, 20 577 iterations is enough to reach satisfactory performance. Using the same loss for inner variational 578 cost and outer training loss (not shown here), see Beauchamp et al. (2022), would require more 579 iterations to converge. Also, constraining the prior to follow the same SPDE simulation ensures 580 to also jointly minimize the OI variational cost (asymptotic convergence of the red line to the 581 yellow star on Fig. 10) which is not the case when looking for an optimal solution within the bi-582 level neural optimization of prior and solver (blue line) that may lead to deviate from the original 583 variational cost to minimize. Let note that by construction, the analytical OI solution is optimal 584 regarding the OI variational cost: it is unbiased with minimal variance. In other words, at a given 585

spatio-temporal location (\mathbf{s}, t) , its variance (which is the local MSE) is minimal. In our case, we compute the global MSE over the entire domain $\mathcal{D} \times [t_k - 2, t_k + 2]$ w.r.t the true state because we have only one single realization to compute this metrics. This is why the global MSE (see \mathcal{L}_1) of the OI may be outperformed by learning-based methods.

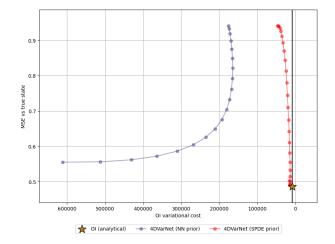


FIG. 10: Optimal Interpolation derived variational cost vs Mean Squared Error (MSE) loss (a) for the gradient-based descent of the variational cost, the classical implementation of the neural scheme and its SPDE-based prior formulation. For the analytical Optimal Interpolation solution (the yellow star), there is no iterations, then a single point is displayed.

590 b. Realistic SSH datasets

Dataset. In this application on Sea Surface Height (SSH) spatio-temporal fields, we focus on a 591 small part of the GULFSTREAM, see Fig. 11, mainly driven by energetic mesoscale dynamics, 592 to illustrate how our framework may help to solve for the oversmoothing of the state-of-the-art 593 Optimal interpolation (OI) and how the SPDE formulation of the prior is a consistent linearization 594 of the dynamics in the data assimilation window that helps to generate ensemble members in the 595 posterior distribution. We use an Observation System Simulation Experiment (OSSE) with the 596 NEMO (Nucleus for European Modeling of the Ocean) model NATL60 high resolution basin-597 scale configuration Molines (2018). Based on this one-year long simulation, we generate pseudo 598 along-track nadir data for the current capabilities of the observation system (Ballarotta et al. 2019) 599 and pseudo wide-swath SWOT data in the context of the upcoming SWOT mission (Metref et al. 600 2020), with additional observation errors (Dufau et al. 2016; Esteban-Fernandez 2014; Gaultier 601 and Ubelmann 2010). The two types of observations may be merged, see Fig. 11, to produce a 602

- one-year long daily datasets of partial and noisy observations of the idealized Ground Truth (GT).
- Last the DUACS operational system (CMEMS/C3S Copernicus program) provides the Optimal
- Interpolation baseline (Taburet et al. 2019) as daily gridded (0.25°x0.25°) products

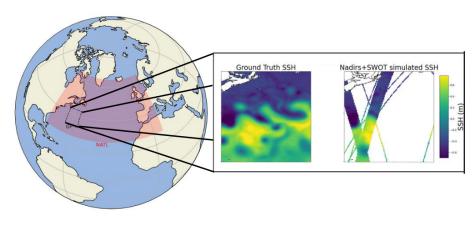


FIG. 11: NATL60 and GULFSTREAM domain and zoom-in picture for one day Ground Truth and accumulated along-track nadir + wide-swath SWOT SSH pseudo-observations

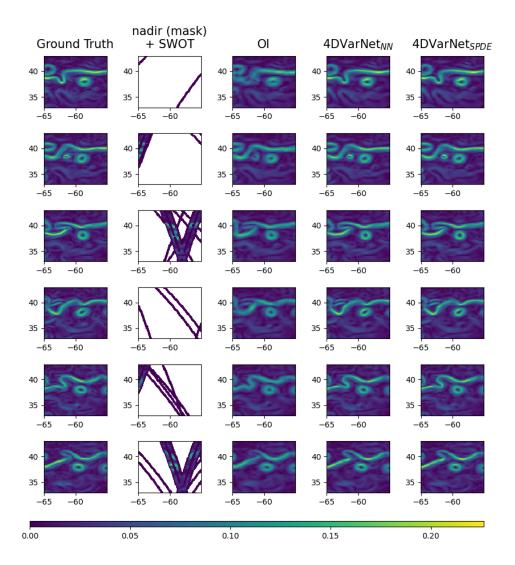


FIG. 12: From left to right: SSH Gradient Ground Truth, Pseudo-observations (nadir mask and SWOT Gradient field), DUACS Optimal Interpolation, and neural variational scheme results obtained with UNet-based and SPDE-based prior operator Φ . A data assimilation window of length 5 is used

Training setting. All the datasets are downscaled from the original resolution of $1/60^{\circ}$ to $1/10^{\circ}$. For the training, the dataset spans from mid-February 2013 to October 2013, while the validation period refers to January 2013. All methods are tested on the test period from October 22, 2012 to December 2, 2012. We still use Adam optimizer with 1100 epochs. Regarding the metrics, we use the ocean data challenge ¹ strategy looking at RMSE-score, and both spatial and temporel

¹https://github.com/ocean-data-challenges/2020a_SSH_mapping_NATL60

⁶¹¹ minimal scales resolved. The reader may refer to (Beauchamp et al. 2023b) for more details.

612

Results. Because we aim at assessing how the use of SPDE priors in the neural architecture is 613 relevant, we used a Gaussian prior formulation $\mathbf{x} \sim \mathcal{N}\left(\mathbf{x}^{b}, \mathbf{Q}_{\theta}^{b}\right)$ where the mean \mathbf{x}^{b} is provided as a 614 first guess by a preliminary 4DVarNet run. Doing so, the SPDE parametrization θ is focused on the 615 reconstruction and surrogate parametrization of the small scales not catched by this deterministic 616 mean. Then, we can reduce the data assimilation window used to estimate the prior mean \mathbf{x}^b , to a 617 reasonable length L = 5 here, so that the storage of multiple (sparse) precision matrices throughout 618 the computational graph remains possible, see Section 3.d. Moving to longer time windows 619 including mesoscale-related autocorrelations (more than 10 days) in this SPDE framework would 620 lead to similar results, see e.g. Febvre et al. (2022) but would require moving to matrix-free 621 formulations, with potential existing solutions, see e.g. Pereira et al. (2022). 622

Last point on this experimental configuration: because the pseudo-observations are subsampled from hourly simulations but we target daily reconstructions, they are noisy due to representativity errors between the two temporal resolutions. This is not currently adressed by our framework where the observation term in the minimization cost, see Eq. 8, is only the L2-norm of the innovations. But it might be easily considered, either by using a known observation error covariance matrix **R** or by learning one of its possible parametrization as an additional feature of the neural scheme.

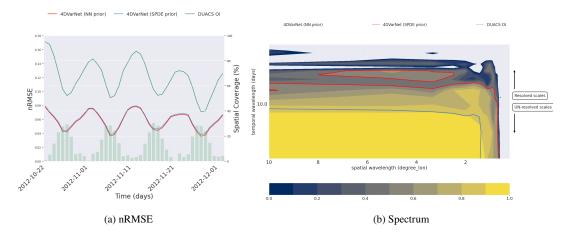


FIG. 13: For DUACS Optimal Interpolation, neural solvers with UNet-based and SPDE-based prior, (a) provides their temporal performance, i.e. nRMSE time series along the BOOST-SWOT DC evaluation period ; and (b) displays their spectral performance, i.e. the PSD-based score is used to evaluate the spatio-temporal scales resolved in the GULFSTREAM domain (yellow area)

As in Example 7.1, we provide in Fig. 12 the reconstructions, as the SSH gradients, obtained 629 from DUACS OI baseline, and both neural solver formulations, the one using as prior a UNet-based 630 parametrization, see e.g. (Beauchamp et al. 2023c) and our SPDE-based formulation. As expected 631 and already seen in previous related studies (Beauchamp et al. 2023b,a; Fablet et al. 2021), the 632 neural schemes improve the baseline by retrieving the dynamics along the main meander of the Gulf 633 Stream and additional small energetic eddies. Again, there is no significant differences between 634 the two neural formulations, which was again expected because the aim of the SPDE formulation 635 is not to improve the mean state estimation obtained when using a neural prior operator, which 636 is even more general, but to provide a a stochastic framework for interpretability and uncertainty 637 quantification. This is supported by Figs. 13a and 13b respectively showing the normalized RMSE 638 and the space-time spectrum along the test period. The periodic improvements of the score are 639 due to the SWOT sampling that does not provide informations every day on this Gulf Stream 640 domain. Overall, the nRMSE is in average improved by 60% when using the neural architecture. 641 For the spectrum, minimal spatio-temporal scales λ_x and λ_t also improve resp. by 30% and 60%. 642 From both figures and scores provided in Table 1, we can see that the UNet formulation of prior 643 Φ leads to a small improvement in the reconstruction, which was expected because this is the 644 only task optimized by the pure neural formulation. Introducing the SPDE formulation leads to 645 optimize both reconstruction and likelihood of the parameters, which is more difficult. Though, the 646 reconstruction performed by the latter is satisfactory enough and very close to the original solution 647 proposed in Fablet et al. (2021). It also competes with other state-of-the-art method available in the 648 ocean data challenges 2020a, among which DUACS OI (Taburet et al. 2019), MIOST (Multi-scale 649 OI) Ardhuin et al. (2020) or a 4DVar scheme based on a QG dynamical model Le Guillou (2022). 650 While DUACS OI has minimal spatial and temporal resolution of 1.22° and 11.15 days, 4DVarNet 651 reaches 0.62° and 4.35 days, which reaches similar order of performance that combining a 4DVar 652 with a QG model. 653

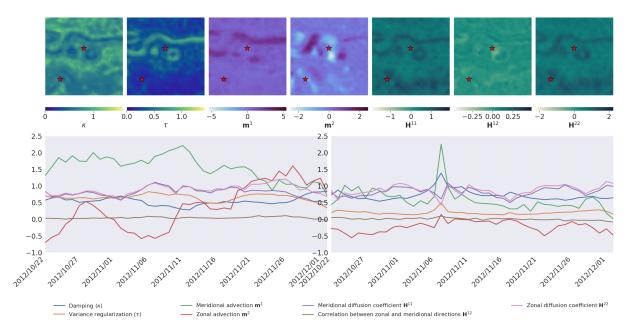


FIG. 14: Parameter estimation of the SPDE prior: in the neural scheme along the 42 days test period and every 6 days. Top, from left to right: τ , \mathbf{m}^1 , \mathbf{m}^2 (advection fields), $\mathbf{H}^{1,1}$, $\mathbf{H}^{1,2}$ and $\mathbf{H}^{2,2}$ (diffusion tensor) estimated on the first day of the test period. Bottom: time series of these parameters along the test period on the two locations identified by the red stars. The first one (left) is located along the Gulf Stream meander and the second one (right) is in the less energetic left-lower part of the domain.

In this very general setup, where the equation-based dataset provide a supervised learning setting 654 on state x, but not on the SPDE parameters θ , two main questions are raised. On the first one: 655 does the parameters retrieved by the iterative solver are interpretable? If considering Fig. 14 that 656 shows parameters τ , \mathbf{m}_1 , \mathbf{m}_2 (advection fields), \mathbf{H}_{11} , \mathbf{H}_{12} and \mathbf{H}_{22} (diffusion tensor) on the first 657 day of the test period, they seem consistent with the SSH field x that partially encodes the SPDE 658 parametrization, which also opens avenue for state-dependent parameters. We also show the time 659 series along the 42 days of the test period for these parameters on two locations of the Gulf Stream 660 domain (red stars): a first one right in the Gulf Stream meander (left time series) and a second 661 one in the left-lower part of the domain, with less variability (right time series). Interestingly, 662 in less energetic areas the parameters are almost perfectly correlated while in the Gulf Stream, 663 they might behave differently. Playing with both damping and variance regularization parameters 664 provide a flexible way to handle complex GP priors with both low and high marginal variances 665 for a given time. This is a key aspect here because the range of possible values attributed to the 666 anomaly generally differs according to the spatio-temporal dynamics of the SSH: it is high along 667

the main meander of the Gulf Stream and eddies not catched up by the OI, and lower elsewhere. This is also supported by generating non-conditional simulations based on these parameters and comparing it to the ground truth anomaly field, see again Fig. 6. The spatio-temporal fields are clearly consistent with the original simulation which also makes the link between our approach and generative modeling. The average of a large number of simulations would be a zero state, while its covariance relates to the model error matrix used in equation-based DA to perturb the members in ensemble methods.

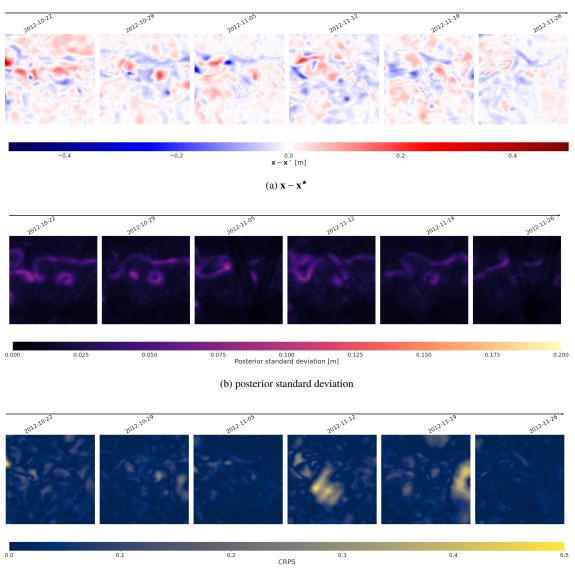
	μ (RMSE)	$\sigma(\text{RMSE})$	λ_x (degree)	λ_t (days)
OI	0.92	0.02	1.22	11.06
MIOST	0.94	0.01	1.18	10.33
4DVar-QG	0.96	0.01	0.66	4.65
4DVarNet (NN prior)	0.96	0.01	0.62	4.35
4DVarNet (SPDE prior)	0.96	0.01	0.64	5.03

TABLE 1: Evaluation of the performance between OI, MIOST, 4DVar-QG, 4DVarNet schemes with UNet and SPDE-based prior parametrization. The OSSE involves 1 SWOT + 4 nadirs as pseudo-observations.

The second question is: does this approach, SPDE-based generation of members, followed by their conditioning with observations, is efficient to estimate the posterior pdf $p_{\mathbf{x}|\mathbf{y}}$? Fig. 15 shows in a) the reconstruction error $\mathbf{x} - \mathbf{x}^*$ for six days along the test period and in b) the empirical posterior standard deviations computed from 200 members. In c), we also provide pointwise Continuous Ranked Probability Score (CRPS) maps to assess the accuracy of the ensemble-based predictions. Given the observations \mathbf{y}_{ijk} , we compute the empirical CDF of the stochastic process *X* at a given spatio-temporal location $\{i, j, k\}$ as $F_{ijk}(z) = \mathbf{P} [X_{ijk} \le z]$:

$$CRPS(F_{ijk},\mathbf{y}_{ijk}) = \int_{-\infty}^{+\infty} \left(F(z) - \mathbf{1}_{z-\mathbf{y}_{ijk}}\right)^2 dz$$

Looking at the estimated standard deviation produced by the ensemble of neural variational reconstructions, they are not as dependent of the observations as an OI scheme would be, which validates the flow-dependency discussed in Section e. We can still see the observation mask as blurry areas ⁶⁸⁵ but the standard deviations are rather increasing and continuous along the main meander of the ⁶⁸⁶ Gulf Stream. The CRPS is often close to zero, which indicates the ensemble of recontruction is ⁶⁸⁷ wholly accurate with a realistic posterior standard deviation. The highest CRPS values observed ⁶⁸⁸ are about 0.4 which remains reasonable, and can be explained by high reconstruction errors outside ⁶⁸⁹ the main Gulf Stream meander that are not correctly handled by the ensemble.



(c) continuous ranked probability score (CRPS) maps

FIG. 15: For six days along the OSSE test period: a) corresponding reconstruction error at the center of the assimilation window, b) ensemble-based posterior standard deviations and c) Continuous Ranked Probability Score

5. Conclusion

We explore a new neural architecture to tackle the reconstruction inverse problem of a dynamical process from partial and potentially noisy observations. We provide a joint end-to-end learning scheme of both stochastic prior models and solvers. The idea is to optimize in the same time the state and the stochastic parametrization of the prior so that we minimize the mean squared error between the reconstruction and the true states, while we are also able to provide uncertainty quantification of the mean state reconstruction, either analytically in the gaussian case, or ensemble-derived in the more general configuration.

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In our work, we draw from recent advances in geostatistical simulations of SPDE-based Gaussian Processes to provide a flexible trainable prior embedded in a neural architecture backboned on variational data assimilation. The SPDE parameters are added as latent variables in an augmented state the trainable solver has to reconstruct. A bi-level optimization scheme is used to optimize in the same time:

• the inner variational cost derived from OI-based formulations, which depends on state \mathbf{x} , observations \mathbf{y} and SPDE parameters $\boldsymbol{\theta}$ and,

• the outer training loss function of the neural architecture, which drives the optimization of the LSTM-based residual solver parameters ω leading to the reconstruction of the augmented state.

The first application of the framework on a diffusion-based GP showed that it reaches the same 709 performance, in terms of MSE w.r.t the ground truth, observed when using a neural-based prior 710 in the classic implementation fo the neural scheme (Fablet et al. 2020), which asymptotically 711 converges towards the optimal solution. In addition, the posterior variance of the mean state 712 derived from the SPDE parameters is close to the true variance of the Optimal Interpolation, 713 which demonstrates the potentiality of the proposed scheme to handle uncertainty quantifications. 714 Indeed, if we showed that retrieving the original set of SPDE parameters might be difficult, the 715 local minimum found after the training process leads to a parametrization with high likelihood, 716 similar diffusion-based spatial patterns and spatio-temporal covariances structures. Though not 717 always physically explainable, the SPDE prior formulation still helps to interpret the dynamical 718

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⁷¹⁹ process in terms of statistical properties.

In a more general setup, we also present an application on Sea Surface Height dynamics based 720 on Observation System Simulation Experiment (OSSE), for which the ground truth is given 721 by a state-of-the-art ocean model and pseudo-observations are generated by a realistic satellite 722 subsampling of the ground truth. In this case, the process is not linear and the gaussian framework 723 does not apply. Then, the idea is to use the GP linear SPDE formulation as a surrogate model to 724 linearize the prior dynamics along the data assimilation window. It provides an efficient way for 725 fast sampling of a huge set of members in the prior distribution within a few minutes. Based on 726 the neural solver, able to handle non-linear and non-gaussian dynamics based on its supervised 727 training, the conditioning of these simulations leads to the estimation of the posterior distribution. 728 The preliminary conclusions made on the GP experiment holds for the reconstruction of the mean 729 state, i.e. no significant differences observed when using a trainable neural prior. The key aspect 730 of the framework is here revealed by the SPDE parametrization, which is fully non-stationary in 731 space and time, and allows for online estimation after training for any new set of input observations, 732 in contrast with other approaches in most of the spatial statistics literature which requires offline 733 new parameter inference. We showed that the prior parametrization is both statistically consistent 734 with the original ground truth used in the training and physically sounded with similar patterns 735 observed in the simulations. Comparing the posterior pdf retrieved from the ensemble members, 736 higher variances are observed in areas of high reconstruction errors which is a good indicator 737 of the framework capability to quantify realistic uncertainties. Additional applications would be 738 necessary on real datasets to complement these preliminary conclusions. 739

740

Regarding the potential extensions of this methodology, it is important to understand that such 741 an SPDE parametrization of the prior term is a way of going to generative modelling, as it is 742 called in the machine learning community. In our case, the SPDE is linear and already provide 743 an efficient way to produce fast, large and realistic ensembles. Promising avenues would be to 744 draw from the existing link between diffusion models and SDE to enrich our framework, provide 745 stochastic non-linear priors and see if and how it helps to improve the results obtained in this work. 746 Also, we made the choice to model the prior with a linear SPDE before conditioning it with our 747 neural solver. A direct use of the SPDE formulation for the posterior would have been possible, 748

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⁷⁴⁹ but restrictive to GP reconstructions which is in most cases rather limited. Back to the non-linear
 ⁷⁵⁰ neural diffusion operators, an other way of addressing the problem would be to directly optimize
 ⁷⁵¹ the sampling in the posterior pdf.

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Data availability statement. The open-source 4DVarNet with SPDE priors version of the 755 code is available at https://github.com/CIA-Oceanix/4dvarnet-starter/tree/maximebeauchamp. 756 The datasets shared through the challenge 2020a also are ocean data avail-757 able on GitHub https://github.com/ocean-data-challenges/2020a_SSH_mapping\ 758 _NATL60, lastaccess: 2022 759

APPENDIX A

761

Finite difference schemes for SPDE

In this work, we propose to use the Euler implicit scheme as a discretization method for the stochastic PDEs. We need to define the discretized version of several differential operators on the 2D regular grid, namely:

765 *a. Discretization of the spatial differential operators*

$$\Delta \mathbf{x}_{i,j} = \frac{\partial^2 \mathbf{x}}{\partial x^2}_{i,j} + \frac{\partial^2 \mathbf{x}}{\partial y^2}_{i,j}$$

$$= \frac{\partial}{\partial x} \left(\frac{\partial \mathbf{x}}{\partial x} \right)_{i,j} + \frac{\partial}{\partial y} \left(\frac{\partial \mathbf{x}}{\partial y} \right)_{i,j}$$

$$= \frac{\partial}{\partial x} \left(\frac{\mathbf{x}_{i+1,j} - \mathbf{x}_{i-1,j}}{2dx} \right) + \frac{\partial}{\partial y} \left(\frac{\mathbf{x}_{i,j+1} - \mathbf{x}_{i,j-1}}{2dy} \right)$$

$$= \left(\frac{\mathbf{x}_{i+1,j}}{dx^2} - \frac{\mathbf{x}_{i,j}}{dx^2} \right) - \left(\frac{\mathbf{x}_{i,j}}{dx^2} - \frac{\mathbf{x}_{i-1,j}}{dx^2} \right)$$

$$+ \left(\frac{\mathbf{x}_{i,j+1}}{dy^2} - \frac{\mathbf{x}_{i,j}}{dy^2} \right) - \left(\frac{\mathbf{x}_{i,j}}{dy^2} - \frac{\mathbf{x}_{i,j-1}}{dy^2} \right)$$

$$= \frac{\mathbf{x}_{i+1,j} - 2\mathbf{x}_{i,j} + \mathbf{x}_{i-1,j}}{dx^2} + \frac{\mathbf{x}_{i,j+1} - 2\mathbf{x}_{i,j} + \mathbf{x}_{i,j-1}}{dy^2}$$
(A1a)

$$\nabla \mathbf{H} \cdot \nabla \mathbf{x}_{i,j} = \left(\mathbf{H}^{1,1} \frac{\partial}{\partial x^2} + \mathbf{H}^{1,2} \frac{\partial}{\partial x \partial y} + \mathbf{H}^{2,1} \frac{\partial}{\partial x \partial y} + \mathbf{H}^{2,2} \frac{\partial}{\partial y^2} \right) \mathbf{x}_{i,j}$$

$$= \mathbf{H}^{1,1} \frac{\mathbf{x}_{i+1,j} - 2\mathbf{x}_{i,j} + \mathbf{x}_{i-1,j}}{dx^2}$$

$$+ \mathbf{H}^{2,2} \frac{\mathbf{x}_{i,j+1} - 2\mathbf{x}_{i,j} + \mathbf{x}_{i,j-1}}{dy^2}$$

$$+ \mathbf{H}^{1,2} \frac{\mathbf{x}_{i+1,j+1} - \mathbf{x}_{i+1,j-1} - \mathbf{x}_{i-1,j+1} + \mathbf{x}_{i-1,j-1}}{4dxdy}$$

$$+ \mathbf{H}^{2,1} \frac{\mathbf{x}_{i+1,j+1} - \mathbf{x}_{i-1,j+1} - \mathbf{x}_{i+1,j-1} + \mathbf{x}_{i-1,j-1}}{4dxdy}$$

$$= \mathbf{H}^{1,1} \frac{\mathbf{x}_{i+1,j} - 2\mathbf{x}_{i,j} + \mathbf{x}_{i-1,j}}{dx^2}$$

$$+ \mathbf{H}^{2,2} \frac{\mathbf{x}_{i,j+1} - 2\mathbf{x}_{i,j} + \mathbf{x}_{i,j-1}}{dy^2}$$

$$+ \mathbf{H}^{1,2} \frac{\mathbf{x}_{i+1,j+1} - \mathbf{x}_{i+1,j-1} - \mathbf{x}_{i-1,j+1} + \mathbf{x}_{i-1,j-1}}{2dxdy}$$
(A1b)

7	6	7

$$\mathbf{m} \cdot \nabla \mathbf{x}_{i,j} = \mathbf{m}^{1} \frac{\partial}{\partial x} \mathbf{x}_{i,j} + \mathbf{m}^{2} \frac{\partial}{\partial y} \mathbf{x}_{i,j}$$
$$= \mathbf{m}^{1} \frac{\mathbf{x}_{i+1,j} - \mathbf{x}_{i-1,j}}{2dx} + \mathbf{m}^{2} \frac{\mathbf{x}_{i,j+1} - \mathbf{x}_{i,j-1}}{2dy}$$
(A1c)

⁷⁶⁸ b. Upwind schemes for advection-dominated SPDE

With such an advection-diffusion framework detailed in Section b, it is known that the solution to 769 the space-centered scheme does not oscillate only when the Peclet number is lower than 2 and the 770 Courant–Friedrichs–Lewy condition (CFL) condition $Cr = dt \left(\frac{\mathbf{m}^2}{dx} + \frac{\mathbf{m}^2}{dy} \right) \le 1$ is satisfied 771 (Lewy et al. 1928; Price et al. 1966). For unsatisfied Peclet conditions, damped oscillations occur 772 with nonreal eigenvalues (Finlayson 1992; Price et al. 1966), while in the limiting case of pure 773 advection $\mathbf{H} \rightarrow \mathbf{0}$, such a scheme would be unconditionally unstable (Finlayson 1992; Strikwerda 774 1989). Because the velocity fiels is allowed to vary in space and time, the CFL number is different 775 at each discrete space-time location (i, j, t). A necessary condition for convergence is that the CFL 776 condition be satisfied at each point location, the velocity and diffusion parameters being unknown, 777

we have to choose the timestep *dt* small enough so that the maximum CFL number (observed in space at each time step) satisfies the CFL condition. Again, because the velocity field is trained, it may happen that the CFL condition is not satisfied if the latter is more and more dominant during the training. A simple way of counteracting this problem is to use an activation function on the two velocity components by clipping their maximum value.

783

One way of addressing this problem of stabilities in FDM when the advection term is predominant 784 over the diffusion relates to the class of upwind schemes (UFDM). It is used to numerically simulate 785 more properly the direction of propagation of the state in a flow field. The first order upwind FDM 786 uses a one-sided finite difference in the upstream direction to approximate the advection term in 787 the transport SPDE. The spatial accuracy of the first-order upwind scheme can be improved by 788 choosing a more accurate finite difference stencil for the approximation of spatial derivative. Let 789 note that UFDM scheme for SPDE eliminates the nonphysical oscillations in the space-centered 790 scheme and generate stable solutions even for very complicated flows. 791

⁷⁹² 1) First-order upwind scheme (UFDM1)

⁷⁹³ Instead of using centered differences:

$$\left(\frac{\partial \mathbf{x}}{\partial x}\right)_{i,j} = \frac{\mathbf{x}_{i+1,j} - \mathbf{x}_{i-1,j}}{2dx}$$
(A2a)

794

$$\left(\frac{\partial \mathbf{x}}{\partial y}\right)_{i,j} = \frac{\mathbf{x}_{i,j+1} - \mathbf{x}_{i,j-1}}{2dy}$$
(A2b)

⁷⁹⁵ We use the one-sided upwind differences :

$$\begin{cases} \left(\frac{\partial \mathbf{x}}{\partial x}\right)_{i,j} = \frac{\mathbf{x}_{i,j} - \mathbf{x}_{i-1,j}}{dx} \text{ if } \mathbf{m}_{i,j}^1 > 0\\ \left(\frac{\partial \mathbf{x}}{\partial x}\right)_{i,j} = \frac{\mathbf{x}_{i+1,j} - \mathbf{x}_{i,j}}{dx} \text{ if } \mathbf{m}_{i,j}^1 < 0 \end{cases}$$
(A3a)

796 and

$$\begin{cases} \left(\frac{\partial \mathbf{x}}{\partial y}\right)_{i,j} = \frac{\mathbf{x}_{i,j} - \mathbf{x}_{i,j-1}}{dy} \text{ if } \mathbf{m}_{i,j}^2 < 0\\ \left(\frac{\partial \mathbf{x}}{\partial y}\right)_{i,j} = \frac{\mathbf{x}_{i,j+1} - \mathbf{x}_{i,j}}{dy} \text{ if } \mathbf{m}_{i,j}^2 < 0 \end{cases}$$
(A3b)

⁷⁹⁷ 2) Third-order upwind scheme (UFDM3)

It can be shown, see e.g. (E. Ewing and Wang 2001) that the UFDM scheme is actually a second-798 order approximation of the SPDE with a modified diffusion term. Along this line, it comes with 799 the family of methods that may introduce excessive numerical diffusion in the solution with large 800 gradients. Thus, we use a third order upwind scheme for the approximation of spatial derivatives 801 with four points instead of two, with only a reduced increase in the degree of sparsity of the 802 precision matrix. This scheme is less diffusive compared to the second-order accurate scheme. It 803 comes with four points instead of two for the approximation, with only a reduced increase in the 804 degree of sparsity of the discretized differential operator. 805

⁸⁰⁶ It can be expressed as follows:

$$\begin{cases} \left(\frac{\partial \mathbf{x}}{\partial x}\right)_{i,j} = \frac{2\mathbf{x}_{i+1,j} + 3\mathbf{x}_{i,j} - 6\mathbf{x}_{i-1,j} + \mathbf{x}_{i-2,j}}{6dx} \text{ if } \mathbf{m}_{i,j}^1 > 0 \\ \left(\frac{\partial \mathbf{x}}{\partial x}\right)_{i,j} = \frac{-\mathbf{x}_{i+2,j} + 6\mathbf{x}_{i+1,j} - 3\mathbf{x}_{i,j} - 2\mathbf{x}_{i-1,j}}{6dx} \text{ if } \mathbf{m}_{i,j}^1 < 0 \end{cases}$$
(A4a)

807 and

$$\left(\left(\frac{\partial \mathbf{x}}{\partial y} \right)_{i,j} = \frac{2\mathbf{x}_{i,j+1} + 3\mathbf{x}_{i,j} - 6\mathbf{x}_{i,j-1} + \mathbf{x}_{i,j-2}}{6dy} \text{ if } \mathbf{m}_{i,j}^2 > 0 \\
\left(\left(\frac{\partial \mathbf{x}}{\partial y} \right)_{i,j} = \frac{-\mathbf{x}_{i,j+2} + 6\mathbf{x}_{i,j+1} - 3\mathbf{x}_{i,j} - 2\mathbf{x}_{i,j-1}}{6dy} \text{ if } \mathbf{m}_{i,j}^2 < 0$$
(A4b)

⁸⁰⁸ c. Discretization of the spatio-temporal SPDE

Based on the discretization of the spatial operators using upwind finite difference schemes for the advection term and centered difference schemes for the diffusion term, we involve an implicit ⁸¹¹ Euler scheme to solve the advection-diffusion SPDE:

$$\frac{\partial \mathbf{x}}{\partial t} + \left\{ \kappa^2(\mathbf{s}, t) + \mathbf{m}(\mathbf{s}, t) \cdot \nabla - \nabla \cdot \mathbf{H}(\mathbf{s}, t) \nabla \right\}^{\alpha/2} \mathbf{x}(\mathbf{s}, t) = \boldsymbol{\tau}(\mathbf{s}, t) \mathbf{z}(\mathbf{s}, t)$$

Because the upwind advection schemes is not symmetric and does not involve the same neighbours in the difference scheme approximation according to the predominant flow direction, we introduce the following generic notations.

$$\begin{cases} \mathbf{m}_{i,j}^{1,t,-} = \left(\frac{\partial \mathbf{x}}{\partial x}\right)_{i,j}^{t} & \text{if } \mathbf{m}_{i,j}^{1,t} > 0 \\ \mathbf{m}_{i,j}^{1,t,+} = \left(\frac{\partial \mathbf{x}}{\partial x}\right)_{i,j}^{t} & \text{if } \mathbf{m}_{i,j}^{1,t} < 0 \end{cases} \quad \text{and} \begin{cases} \mathbf{m}_{i,j}^{2,t,-} = \left(\frac{\partial \mathbf{x}}{\partial y}\right)_{i,j}^{t} & \text{if } \mathbf{m}_{i,j}^{2,t} > 0 \\ \mathbf{m}_{i,j}^{2,t,+} = \left(\frac{\partial \mathbf{x}}{\partial y}\right)_{i,j}^{t} & \text{if } \mathbf{m}_{i,j}^{2,t} < 0 \end{cases}$$

and by denoting $\mathbf{a}_{i,j}^{1,t,+} = \max(\mathbf{m}_{i,j}^{1,t}, 0)$, $\mathbf{a}_{i,j}^{1,t,-} = \min(\mathbf{m}_{i,j}^{1,t}, 0)$, $\mathbf{a}_{i,j}^{2,t,+} = \max(\mathbf{m}_{i,j}^{2,t}, 0)$, $\mathbf{a}_{i,j}^{2,t,-} = \min(\mathbf{m}_{i,j}^{2,t}, 0)$, the resulting UFDM, whatever the order of the scheme, can be written in its compact form as:

$$\begin{aligned} \mathbf{x}_{i,j}^{t+1} &= \mathbf{x}_{i,j}^{t} + dt \left[\kappa_{i,j}^{t} \mathbf{x}_{i,j}^{t} + \left(\mathbf{a}_{i,j}^{1,t,+} \mathbf{m}_{i,j}^{1,t,-} + \mathbf{a}_{i,j}^{1,t,-} \mathbf{m}_{i,j}^{1,t,+} \right) + \left(\mathbf{a}_{i,j}^{2,t,+} \mathbf{m}_{i,j}^{2,t,-} + \mathbf{a}_{i,j}^{2,t,-} \mathbf{m}_{i,j}^{2,t,+} \right) \\ &+ \mathbf{H}_{i,j}^{1,1,t} \frac{\mathbf{x}_{i+1,j}^{t} - 2\mathbf{x}_{i,j}^{t} + \mathbf{x}_{i-1,j}^{t}}{dx^{2}} + \mathbf{H}_{i,j}^{2,2,t} \frac{\mathbf{x}_{i,j+1}^{t} - 2\mathbf{x}_{i,j}^{t} + \mathbf{x}_{i,j-1}^{t}}{dy^{2}} \\ &+ \mathbf{H}_{i,j}^{1,2,t} \frac{\mathbf{x}_{i+1,j+1}^{t} - \mathbf{x}_{i+1,j-1}^{t} - \mathbf{x}_{i-1,j+1}^{t} + \mathbf{x}_{i-1,j-1}^{t} + \mathbf{t}_{i,j}^{t} \mathbf{z}_{i,j} \right] \end{aligned}$$

Using again notation $i = \lfloor k/N_x \rfloor$ and $j = k \mod N_x$, operator A_t associated to UFDM1 finally writes:

$$\mathbf{A}_{i,j}^{1,2,t}/2dxdy \text{ if } l = k \pm (N_x + 1)$$

$$\mathbf{H}_{i,j}^{1,2,t}/2dxdy \text{ if } l = k \pm (N_x - 1)$$

$$-\mathbf{H}_{i,j}^{1,1,t}/dx^2 + \left(\mathbf{a}_{i,j}^{1,t,+} + \mathbf{a}_{i,j}^{1,t,-}\right)/dx \text{ if } l = k - 1$$

$$-\mathbf{H}_{i,j}^{1,1,t}/dx^2 + \left(\mathbf{a}_{i,j}^{1,t,+} + \mathbf{a}_{i,j}^{1,t,-}\right)/dx \text{ if } l = k + 1$$

$$-\mathbf{H}_{i,j}^{2,2,t}/dy^2 + \left(\mathbf{a}_{i,j}^{2,t,+} + \mathbf{a}_{i,j}^{1,t,-}\right)/dy \text{ if } l = k - N_x \qquad (A5)$$

$$-\mathbf{H}_{i,j}^{2,2,t}/dy^2 + \left(\mathbf{a}_{i,j}^{2,t,+} + \mathbf{a}_{i,j}^{1,t,-}\right)/dy \text{ if } l = k + N_x$$

$$\left(\kappa_{i,j}^t\right)^2 + 2(\mathbf{H}_{i,j}^{1,1,t}/dx^2 + \mathbf{H}_{i,j}^{2,2,t}/dy^2)$$

$$+ \left(\mathbf{a}_{i,j}^{1,t,+} + \mathbf{a}_{i,j}^{1,t,-}\right)/dx + \left(\mathbf{a}_{i,j}^{2,t,+} + \mathbf{a}_{i,j}^{2,t,-}\right)/dy \text{ if } k = l$$

$$0 \text{ otherwise}$$

⁸²⁰ The same operator associated to UFDM3 writes:

$$\begin{split} \mathbf{A}_{k,l}(t) &= \\ \left\{ \begin{aligned} \mathbf{H}_{l,j}^{1,2,l}/2dxdy & \text{if } l = k \pm (N_x + 1) \\ \mathbf{H}_{l,j}^{1,2,l}/2dxdy & \text{if } l = k \pm (N_x - 1) \\ \mathbf{a}_{l,j}^{1,1,l}/6dx & \text{if } l = k - 2 \\ -\mathbf{H}_{l,j}^{1,1,l}/dx^2 - \left(6\mathbf{a}_{l,j}^{1,l,+} + 2\mathbf{a}_{l,j}^{1,l,-} \right) / 6dx & \text{if } l = k - 1 \\ -\mathbf{H}_{l,j}^{1,1,l}/dx^2 + \left(2\mathbf{a}_{l,j}^{1,l,+} + 6\mathbf{a}_{l,j}^{1,l,-} \right) / 6dx & \text{if } l = k + 1 \\ -\mathbf{a}_{l,j}^{1,l,-} / 6dx & \text{if } l = k + 2 \\ \left\{ \mathbf{a}_{l,j}^{2,l,+} / 6dy & \text{if } l = k - 2Nx & (A6) \\ -\mathbf{H}_{l,j}^{2,2,l}/dy^2 - \left(6\mathbf{a}_{l,j}^{2,l,+} + 2\mathbf{a}_{l,j}^{2,l,-} \right) / 6dy & \text{if } l = k - N_x \\ -\mathbf{H}_{l,j}^{2,l,l} / dy^2 + \left(2\mathbf{a}_{l,j}^{2,l,+} + 6\mathbf{a}_{l,j}^{2,l,-} \right) / 6dy & \text{if } l = k + N_x \\ -\mathbf{a}_{l,j}^{2,l,-} / 6dy & \text{if } l = k + 2Nx \\ \left(k_{l,j}^{t} \right)^2 + 2(\mathbf{H}_{l,j}^{1,l,t} / dx^2 + \mathbf{H}_{l,j}^{2,2,l} / dy^2) \\ + 3\left(\mathbf{a}_{l,j}^{1,l,+} - \mathbf{a}_{l,j}^{1,l,-} \right) / 6dx \\ + 3\left(\mathbf{a}_{l,j}^{2,l,+} - \mathbf{a}_{l,j}^{2,l,-} \right) / 6dy & \text{if } k = l \\ \end{split} \right.$$

and 0 otherwise.

APPENDIX B

823

SPDE-based precision matrix

Let define the spatio-temporal SPDE prior $\mathbf{x} = {\mathbf{x}_0, \dots, \mathbf{x}_{Ldt}}$. From now on, $\mathbf{x}_0 \sim \mathcal{N}(\mathbf{0}, \mathbf{P}_0)$ denotes the initial state and $\mathbf{Q}_0 = \mathbf{P}_0^{-1}$ is always taken as the precision matrix obtained after a given stabilization run, i.e. the evolution of the dynamical system over N timesteps using as stationary parameters the initial parametrization $\boldsymbol{\theta}_0$ of the SPDE at time t = 0. We can rewrite :

$$\{\mathbf{x}_0,\cdots,\mathbf{x}_{Ldt}\} = \mathbf{M}_G \begin{bmatrix} \mathbf{x}_0 \\ \mathbf{z} \end{bmatrix}$$

with $\mathbf{z} = [\mathbf{z}_1, ..., \mathbf{z}_t]^T$ and

$$\mathbf{M}_{G} = \begin{bmatrix} \mathbf{I} & 0 & 0 & 0 & 0 & \dots & 0 \\ \mathbf{M}_{1} & \mathbf{T}_{1} & 0 & 0 & 0 & \dots & 0 \\ \mathbf{M}_{2}\mathbf{M}_{1} & \mathbf{M}_{2}\mathbf{T}_{1} & \mathbf{T}_{2} & 0 & 0 & \dots & 0 \\ \mathbf{M}_{3}\mathbf{M}_{2}\mathbf{M}_{1} & \mathbf{M}_{3}\mathbf{M}_{2}\mathbf{T}_{1} & \mathbf{M}_{3}\mathbf{T}_{2} & \mathbf{T}_{3} & 0 & \dots & 0 \\ \vdots & \ddots & \ddots & \ddots & \ddots & \ddots & 0 \\ \vdots & \ddots & \ddots & \ddots & \ddots & \ddots & \ddots & 0 \\ \vdots & \ddots & \ddots & \ddots & \ddots & \ddots & \ddots & \mathbf{T}_{L} \end{bmatrix}$$

⁸²⁹ Despite its apparent complexity, \mathbf{M}_G has a particular structure which allows to easily compute ⁸³⁰ its inverse:

$$\mathbf{M}_{G}^{-1} = \begin{bmatrix} \mathbf{I} & \mathbf{0} & \mathbf{0} & \mathbf{0} & \mathbf{0} & \dots & \mathbf{0} \\ -\mathbf{T}_{1}^{-1}\mathbf{M}_{1} & \mathbf{T}_{1}^{-1} & \mathbf{0} & \mathbf{0} & \mathbf{0} & \dots & \mathbf{0} \\ \mathbf{0} & -\mathbf{T}_{2}^{-1}\mathbf{M}_{2} & \mathbf{T}_{2}^{-1} & \mathbf{0} & \mathbf{0} & \dots & \mathbf{0} \\ \mathbf{0} & \mathbf{0} & -\mathbf{T}_{3}^{-1}\mathbf{M}_{3} & \mathbf{T}_{3}^{-1} & \mathbf{0} & \dots & \mathbf{0} \\ \mathbf{0} & \ddots & \ddots & \ddots & \ddots & \ddots & \mathbf{0} \\ \vdots & \ddots & \ddots & \ddots & \ddots & \ddots & \mathbf{0} \\ \mathbf{0} & \ddots & \ddots & \ddots & \ddots & \ddots & \mathbf{0} \\ \mathbf{0} & \ddots & \ddots & \ddots & \ddots & \mathbf{0} \\ \mathbf{0} & \ddots & \ddots & \ddots & \ddots & \mathbf{1} \\ \mathbf{1} \mathbf{M}_{L} & \mathbf{T}_{L}^{-1} \end{bmatrix}$$

All this information is embedded in \mathbf{Q}^b (Eq. B2), which is the inverse of the prior covariance matrix **B**:

$$\mathbf{Q}^{b} = \mathbf{B}^{-1} = \begin{bmatrix} \mathbf{P}_{0} & \mathbf{P}_{0,1} & \dots & \dots & \mathbf{P}_{0,L} \\ \mathbf{P}_{1,0} & \mathbf{P}_{1} & \dots & \dots & \mathbf{P}_{1,L} \\ \vdots & \mathbf{P}_{2,1} & \mathbf{P}_{2} & \dots & \mathbf{P}_{2,L} \\ \vdots & \ddots & \ddots & \ddots & \vdots \\ \mathbf{P}_{L-1,0} & \dots & \dots & \mathbf{P}_{L-1} & \mathbf{P}_{L-1,L} \end{bmatrix}^{-1} = \mathbf{M}_{G}^{-1} \begin{bmatrix} \mathbf{P}_{0}^{-1} & 0 & \dots & 0 \\ 0 & \mathbf{I} & \dots & 0 \\ \vdots & \ddots & \ddots & \vdots \\ 0 & 0 & \dots & \mathbf{I} \end{bmatrix} \mathbf{M}_{G}^{-1}$$
(B1)

By denoting $\mathbf{S}_k = \mathbf{T}_k \mathbf{T}_k^{\mathrm{T}}$, we have

$$\mathbf{Q}^{b} = \begin{bmatrix} \mathbf{P}_{0}^{-1} + \mathbf{M}_{1}^{T} \mathbf{S}_{1}^{-1} \mathbf{M}_{1} & -\mathbf{M}_{1}^{T} \mathbf{S}_{2}^{-1} \mathbf{M}_{2} & -\mathbf{M}_{2}^{T} \mathbf{S}_{2}^{-1} & 0 & 0 & \dots & 0 \\ -\mathbf{S}_{1}^{-1} \mathbf{M}_{1} & \mathbf{S}_{1}^{-1} + \mathbf{M}_{2}^{T} \mathbf{S}_{2}^{-1} \mathbf{M}_{2} & -\mathbf{M}_{2}^{T} \mathbf{S}_{2}^{-1} & 0 & 0 & \dots & 0 \\ 0 & -\mathbf{S}_{2}^{-1} \mathbf{M}_{2} & \mathbf{S}_{2}^{-1} + \mathbf{M}_{3}^{T} \mathbf{S}_{3}^{-1} \mathbf{M}_{3} - \mathbf{M}_{3}^{T} \mathbf{S}_{3}^{-1} & 0 & \dots & 0 \\ 0 & -\mathbf{S}_{2}^{-1} \mathbf{M}_{2} & \mathbf{S}_{2}^{-1} + \mathbf{M}_{3}^{T} \mathbf{S}_{3}^{-1} \mathbf{M}_{3} - \mathbf{M}_{3}^{T} \mathbf{S}_{3}^{-1} & 0 & \dots & 0 \\ 0 & \mathbf{N}_{1} & \mathbf{N}_{2} & \mathbf{N}_{2} & \mathbf{N}_{2} & \mathbf{N}_{2} & \mathbf{N}_{3} - \mathbf{N}_{3}^{T} \mathbf{S}_{3}^{-1} \mathbf{N}_{3} - \mathbf{M}_{3}^{T} \mathbf{S}_{3}^{-1} \mathbf{N}_{3} - \mathbf{M}_{3}^{T} \mathbf{S}_{3}^{-1} \mathbf{N}_{3} - \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} - \mathbf{N}_{3}^{T} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} - \mathbf{N}_{3}^{T} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} - \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3}^{T} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3}^{T} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3}^{T} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3}^{T} \mathbf{N}_{3}^{T} \mathbf{N}_{3}^{T} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3}^{T} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3} \mathbf{N}_{3}^{T} \mathbf{N}_{3} \mathbf{N}_{3}^{T} \mathbf{$$

Because of the formulation of \mathbf{M}_t and \mathbf{T}_t , the precision matrix \mathbf{Q}^b with the FDM scheme also writes:

$$\mathbf{Q}^{b} = \frac{1}{dt} \begin{bmatrix} \mathbf{P}_{0}^{-1} + \widetilde{\mathbf{Q}}_{s,1} & -\widetilde{\mathbf{Q}}_{s,1} \mathbf{M}_{1}^{-1} & 0 & 0 & 0 & \dots & 0 \\ -(\mathbf{M}_{1}^{-1})^{-1} \widetilde{\mathbf{Q}}_{s,1} \mathbf{M}_{1}^{-1} \widetilde{\mathbf{Q}}_{s,2} \mathbf{M}_{2}^{-1} \widetilde{\mathbf{Q}}_{s,2} \mathbf{M}_{2}^{-1} & 0 & 0 & \dots & 0 \\ 0 & -(\mathbf{M}_{2}^{-1})^{-1} \widetilde{\mathbf{Q}}_{s,2} \mathbf{M}_{2}^{-1} \widetilde{\mathbf{Q}}_{s,2} \mathbf{M}_{2}^{-1} \widetilde{\mathbf{Q}}_{s,3} - \widetilde{\mathbf{Q}}_{s,3} \mathbf{M}_{3}^{-1} & 0 & \dots & 0 \\ 0 & -(\mathbf{M}_{2}^{-1})^{-1} \widetilde{\mathbf{Q}}_{s,2} \mathbf{M}_{2}^{-1} \widetilde{\mathbf{Q}}_{s,2} \mathbf{M}_{2}^{-1} \widetilde{\mathbf{Q}}_{s,3} - \widetilde{\mathbf{Q}}_{s,3} \mathbf{M}_{3}^{-1} & 0 & \dots & 0 \\ 0 & \ddots & \ddots & \ddots & \ddots & \ddots & \ddots & 0 \\ 0 & \ddots & 0 \\ \vdots & \ddots & 0 \\ \vdots & \ddots & \ddots & \ddots & \ddots & \ddots & \ddots & 0 \\ 0 & \ddots & \ddots & \ddots & 0 & -(\mathbf{M}_{L-1}^{-1})^{-1} \widetilde{\mathbf{Q}}_{s,L-1} \mathbf{M}_{L}^{-1} \widetilde{\mathbf{Q}}_{s,L} - \mathbf{M}_{L}^{-1} \widetilde{\mathbf{Q}}_{s,L} \mathbf{M}_{L} \end{bmatrix} \\ 0 & \ddots & \ddots & \ddots & 0 & -(\mathbf{M}_{L}^{-1})^{-1} \widetilde{\mathbf{Q}}_{s,L} \mathbf{M}_{L}^{-1} \widetilde{\mathbf{Q}}_{s,L} \mathbf{M}_{L}^{-1} \mathbf{M}_{L}^{-1} \widetilde{\mathbf{Q}}_{s,L} \mathbf{M}_{L} \end{bmatrix}$$
(B3)

where $\widetilde{\mathbf{Q}}_{s,t}$ is the precision matrix of the colored noise weighted by the non-uniform regularization variance τ_t .

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APPENDIX C

PyTorch implementation of sparse linear algebra
 Currently, two pieces of codes are missing in the PyTorch sparse linear algebra to achieve a fully
 sparse implementation of our algorithm: the automatic differentiation tools for

• solving sparse linear systems

• running sparse Cholesky decomposition

First, regarding the implementation of the backward pass for solving linear systems, we start by writing the forward pass of this system of equation:

$$x = \mathbf{A}^{-1}\mathbf{b}$$

where **A** denotes a 2*D* square matrix and **b** a one-dimensional vector.

⁸⁴⁸ We need to provide the gradients wrt both **A** and **b**:

$$\frac{\partial L}{\partial \mathbf{b}} = \frac{\partial L}{\partial x_i} \frac{\partial x_i}{\partial \mathbf{b}_j} = \frac{\partial L}{\partial x_i} \frac{\partial}{\partial \mathbf{b}_k} (\mathbf{A}_{ij}^{-1} \mathbf{b}_j) = \frac{\partial L}{\partial x_i} \mathbf{A}_{ij}^{-1} \frac{\partial \mathbf{b}_j}{\partial \mathbf{b}_k}$$
$$= \frac{\partial L}{\partial x_i} \mathbf{A}_{ij}^{-1} \delta_{jk} = \frac{\partial L}{\partial x_i} \mathbf{A}_{ik}^{-1} = (\mathbf{A}^{-1})^T \frac{\partial L}{\partial x}$$
$$= \text{solve}(\mathbf{A}^T, \frac{\partial L}{\partial x})$$

849 and

$$\frac{\partial L}{\partial \mathbf{A}} = \frac{\partial L}{\partial x_i} \frac{\partial x_i}{\partial \mathbf{A}_{mn}} = \frac{\partial L}{\partial x_i} \frac{\partial}{\partial \mathbf{A}_{mn}} (\mathbf{A}_{ij}^{-1} \mathbf{b}_j)$$
$$= -\frac{\partial L}{\partial x_i} \mathbf{A}_{ij}^{-1} \frac{\partial \mathbf{A}_{jk}}{\partial \mathbf{A}_{mn}} \mathbf{A}_{kl}^{-1} \mathbf{b}_l$$
$$= -\frac{\partial L}{\partial x_i} \mathbf{A}_{ij}^{-1} \delta_{jm} \delta_{kn} \mathbf{A}_{kl}^{-1} \mathbf{b}_l$$
$$= -\frac{\partial L}{\partial x_i} \mathbf{A}_{im}^{-1} \mathbf{A}_{nl}^{-1} \mathbf{b}_l$$
$$= -\left(\left((\mathbf{A}^{-1})^T \frac{\partial L}{\partial x} \right) \otimes \left(\mathbf{A}^{-1} \mathbf{b} \right) \right)$$
$$= -\frac{\partial L}{\partial \mathbf{b}} \otimes x$$

where we used the einstein summation convention during the derivations, as well as the following
 identity:

$$\frac{\partial (\mathbf{A}^{-1})}{\partial p} = -\mathbf{A}^{-1} \frac{\partial \mathbf{A}}{\partial p} \mathbf{A}^{-1}.$$

These two expressions $\frac{\partial L}{\partial \mathbf{b}}$ and $\frac{\partial L}{\partial \mathbf{A}}$ are easily implemented in PyTorch based on sparse representations.

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Second, the backward pass for a sparse Cholesky decomposition may be found for instance in Seeger et al. (2019): given a symmetric, positive definite matrix **A**, its Cholesky factor **L** is lower triangular with positive diagonal, such that the forward pass of the Cholesky decomposition is defined as $\mathbf{A} = \mathbf{L}\mathbf{L}^{T}$. Given the output gradient $\overline{\mathbf{L}}$ and the Cholesky factor **L**, the backward pass compute the input gradient $\overline{\mathbf{A}}$ defined as :

$$\overline{\mathbf{A}} = \frac{1}{2} \mathbf{L}^{-\mathrm{T}} lt u (\mathbf{L}^{\mathrm{T}} \overline{\mathbf{L}}) \mathbf{L}^{-1}$$
(C1)

where $ltu(\mathbf{X})$ generates a symmetric matrix by copying the lower triangle to the upper triangle. In our work, such an operation is useful when optimizing the likelihood in the outer training cost function of the neural scheme, see Eq. 20, which involves the determinant of sparse matrices Cholesky decomposition.

References

M. Andrychowicz, M. Denil, S. Gomez, M. W. Hoffman, D. Pfau, T. Schaul, B. Shillingford, and
 N. De Freitas. Learning to learn by gradient descent by gradient descent. In *Advances in neural information processing systems*, pages 3981–3989, 2016.

F. Ardhuin, C. Ubelmann, G. Dibarboure, L. Gaultier, A. Ponte, M. Ballarotta, and Y. Faugère.

Reconstructing ocean surface current combining altimetry and future spaceborne doppler data.

Earth and Space Science Open Archive, page 22, 2020. doi: 10.1002/essoar.10505014.1. URL

https://doi.org/10.1002/essoar.10505014.1.

L. Arras, J. Arjona-Medina, M. Widrich, G. Montavon, M. Gillhofer, K.-R. Müller, S. Hochreiter,
and W. Samek. Explaining and interpreting lstms. *Lecture Notes in Computer Science*, pages
211–238, 2019. ISSN 1611-3349. doi: 10.1007/978-3-030-28954-6_11. URL http://dx.
doi.org/10.1007/978-3-030-28954-6_11.

M. Asch, M. Bocquet, and M. Nodet. *Data Assimilation*. Fundamentals of Algorithms. Society
 for Industrial and Applied Mathematics, Dec. 2016. ISBN 978-1-61197-453-9. doi: 10.1137/1.
 9781611974546. URL https://doi.org/10.1137/1.9781611974546.

M. Ballarotta, C. Ubelmann, M.-I. Pujol, G. Taburet, F. Fournier, J.-F. Legeais, Y. Faugère,
 A. Delepoulle, D. Chelton, G. Dibarboure, and N. Picot. On the resolutions of ocean altimetry
 maps. *Ocean Science*, 15(4):1091–1109, 2019. doi: 10.5194/os-15-1091-2019. URL https:
 //www.ocean-sci.net/15/1091/2019/.

M. Beauchamp, J. Thompson, H. Georgenthum, Q. Febvre, and R. Fablet. Learning neural optimal
 interpolation models and solvers, 2022. URL https://arxiv.org/abs/2211.07209.

M. Beauchamp, Q. Febvre, and R. Fablet. Ensemble-based 4dvarnet uncertacinty quantification for
 the reconstruction of sea surface height dynamics. *Environmental Data Science*, 2:e18, 2023a.
 doi: 10.1017/eds.2023.19.

M. Beauchamp, Q. Febvre, H. Georgenthum, and R. Fablet. 4dvarnet-ssh: end-to-end learning
 of variational interpolation schemes for nadir and wide-swath satellite altimetry. *Geoscientific Model Development*, 16(8):2119–2147, 2023b. doi: 10.5194/gmd-16-2119-2023. URL https:
 //gmd.copernicus.org/articles/16/2119/2023/.

M. Beauchamp, Q. Febvre, J. Thompson, H. Georgenthum, and R. Fablet. Learning neural
 optimal interpolation models and solvers. In J. Mikyška, C. de Mulatier, M. Paszynski, V. V.
 Krzhizhanovskaya, J. J. Dongarra, and P. M. Sloot, editors, *Computational Science – ICCS 2023*,
 pages 367–381, Cham, 2023c. Springer Nature Switzerland. ISBN 978-3-031-36027-5.

D. Bolin and J. Wallin. Spatially adaptive covariance tapering. *Spatial Statistics*, 18:163–178,
 2016. ISSN 2211-6753. doi: https://doi.org/10.1016/j.spasta.2016.03.003. URL https://
 www.sciencedirect.com/science/article/pii/S2211675316000245. Spatial Statistics
 Avignon: Emerging Patterns.

P. Boudier, A. Fillion, S. Gratton, S. Gürol, and S. Zhang. Data assimilation networks. *Journal of Advances in Modeling Earth Systems*, 15(4):e2022MS003353, 2023. doi: https://doi.org/
 10.1029/2022MS003353. URL https://agupubs.onlinelibrary.wiley.com/doi/abs/
 10.1029/2022MS003353. e2022MS003353 2022MS003353.

J. Brajard, A. Carrassi, M. Bocquet, and L. Bertino. Combining data assimilation and machine learning to infer unresolved scale parametrization, 2021. URL https:// royalsocietypublishing.org/doi/abs/10.1098/rsta.2020.0086.

A. Carrassi, M. Bocquet, L. Bertino, and G. Evensen. Data assimilation in the geosciences: An
 overview of methods, issues, and perspectives. *WIREs Climate Change*, 9(5):e535, 2018. doi:
 https://doi.org/10.1002/wcc.535. URLhttps://onlinelibrary.wiley.com/doi/abs/10.
 1002/wcc.535.

Y. Chen, T. Davis, W. Hager, and S. Rajamanickam. Algorithm 887: Cholmod, supernodal sparse
cholesky factorization and update/downdate. *ACM Trans. Math. Softw.*, 35, 01 2008.

S. Cheng and M. Qiu. Observation error covariance specification in dynamical systems for data assimilation using recurrent neural networks. *Neural Computing and Applications*, 34(16):
13149–13167, Aug 2022. ISSN 1433-3058. doi: 10.1007/s00521-021-06739-4. URL https://doi.org/10.1007/s00521-021-06739-4.

S. Cheng, C. Quilodran-Casas, S. Ouala, A. Farchi, C. Liu, P. Tandeo, R. Fablet, D. Lucor, B. Iooss,
 J. Brajard, D. Xiao, T. Janjic, W. Ding, Y. Guo, A. Carrassi, M. Bocquet, and R. Arcucci. Machine

- learning with data assimilation and uncertainty quantification for dynamical systems: a review,
 2023.
- J. Chilès and P. Delfiner. *Geostatistics : modeling spatial uncertainty*. Wiley, New-York, second edition, 2012.
- J.-P. Chilès and N. Desassis. *Fifty Years of Kriging*, pages 589–612. Springer International
 Publishing, Cham, 2018. ISBN 978-3-319-78999-6. doi: 10.1007/978-3-319-78999-6_29.
 URL https://doi.org/10.1007/978-3-319-78999-6_29.
- L. Clarotto, D. Allard, T. Romary, and N. Desassis. The spde approach for spatio-temporal datasets
 with advection and diffusion, 2022. URL https://arxiv.org/abs/2208.14015.
- F. Counillon and L. Bertino. Ensemble optimal interpolation: multivariate properties in the gulf of
 mexico. *Tellus A*, 61(2):296–308, 2009. doi: https://doi.org/10.1111/j.1600-0870.2008.00383.
- x. URL https://onlinelibrary.wiley.com/doi/abs/10.1111/j.1600-0870.2008. 00383.x.
- N. Cressie and C. Wikle. *Statistics for Spatio-Temporal Data*. John Wiley & Sons, 2015. ISBN 978-1-119-24304-5.
- L. Dinh, J. Sohl-Dickstein, and S. Bengio. Density estimation using real nvp, 2017.

 A. Doury, S. Somot, S. Gadat, A. Ribes, and L. Corre. Regional climate model emulator based on deep learning: concept and first evaluation of a novel hybrid downscaling approach. *Climate Dynamics*, 60(5):1751–1779, Mar 2023. ISSN 1432-0894. doi: 10.1007/s00382-022-06343-9.
 URL https://doi.org/10.1007/s00382-022-06343-9.

C. Dufau, M. Orsztynowicz, G. Dibarboure, R. Morrow, and P.-Y. Le Traon. Mesoscale resolution
 capability of altimetry: Present and future. *Journal of Geophysical Research: Oceans*, 121(7):
 4910–4927, 2016. doi: 10.1002/2015JC010904. URL https://agupubs.onlinelibrary.
 wiley.com/doi/abs/10.1002/2015JC010904.

R. E. Ewing and H. Wang. A summary of numerical methods for time-dependent advectiondominated partial differential equations. *Journal of Computational and Applied Mathematics*,
128(1):423–445, 2001. ISSN 0377-0427. doi: https://doi.org/10.1016/S0377-0427(00)00522-7.

- ⁹⁴⁶ URL https://www.sciencedirect.com/science/article/pii/S0377042700005227.
 ⁹⁴⁷ Numerical Analysis 2000. Vol. VII: Partial Differential Equations.
- D. Esteban-Fernandez. Swot project mission performance and error budget document. Technical
 report, JPL, NASA, 2014.
- ⁹⁵⁰ G. Evensen. *Data Assimilation*. Springer Berlin Heidelberg, Berlin, Heidelberg, 2009.
 ⁹⁵¹ ISBN 9783642037108 9783642037115. URL http://link.springer.com/10.1007/
 ⁹⁵² 978-3-642-03711-5.
- ⁹⁵³ G. Evensen, , F. C. Vossepoel, and P. J. van Leeuwen. *Data Assimilation Fundamentals*. Springer
 ⁹⁵⁴ Textbooks in Earth Sciences, Geography and Environment (STEGE). Springer Cham, 2022.
 ⁹⁵⁵ ISBN 978-3-030-96708-6. doi: 10.1007/978-3-030-96709-3. URL https://doi.org/10.
 ⁹⁵⁶ 1007/978-3-030-96709-3.
- ⁹⁵⁷ R. Fablet, L. Drumetz, and F. Rousseau. Joint learning of variational representations and solvers
 ⁹⁵⁸ for inverse problems with partially-observed data. *arXiv:2006.03653 [physics]*, 2020. URL
 ⁹⁵⁹ https://arxiv.org/abs/2006.03653. arXiv:2006.03653.
- R. Fablet, B. Chapron, L. Drumetz, E. Mémin, O. Pannekoucke, and F. Rousseau.
 Learning variational data assimilation models and solvers. *Journal of Advances in Modeling Earth Systems*, 13(10):e2021MS002572, 2021. doi: https://doi.org/10.1029/2021MS002572. URL https://agupubs.onlinelibrary.wiley.com/doi/abs/10.
 1029/2021MS002572. e2021MS002572 2021MS002572.
- Q. Febvre, R. Fablet, J. L. Sommer, and C. Ubelmann. Joint calibration and mapping of satellite
 altimetry data using trainable variational models. In *ICASSP 2022 2022 IEEE International Conference on Acoustics, Speech and Signal Processing (ICASSP)*, pages 1536–1540, 2022. doi:
- ⁹⁶⁸ 10.1109/ICASSP43922.2022.9746889.
- B. A. Finlayson. *Numerical Methods for Problems with Moving Fronts*. Ravenna Park Publishing,
 Seattle, 1992.
- ⁹⁷¹ G.-A. Fuglstad, F. Lindgren, D. Simpson, and H. Rue. Exploring a new class of non-stationary
 ⁹⁷² spatial gaussian random fields with varying local anisotropy. *Statistica Sinica*, 25(1):115–133,
 ⁹⁷³ 2015a. ISSN 1017-0405. doi: 10.5705/ss.2013.106w.

G.-A. Fuglstad, D. Simpson, F. Lindgren, and H. Rue. Does non-stationary spatial data always
 require non-stationary random fields? *Spatial Statistics*, 14:505–531, 2015b. ISSN 2211-6753.
 doi: https://doi.org/10.1016/j.spasta.2015.10.001. URL https://www.sciencedirect.com/
 science/article/pii/S2211675315000780.

G.-A. Fuglstad, D. Simpson, F. Lindgren, and H. Rue. Does non-stationary spatial data always
 require non-stationary random fields? *Spatial Statistics*, 14:505 – 531, 2015c. ISSN 2211-6753.
 doi: https://doi.org/10.1016/j.spasta.2015.10.001. URL http://www.sciencedirect.com/
 science/article/pii/S2211675315000780.

⁹⁸² R. Furrer, M. G. Genton, and D. Nychka. Covariance tapering for interpolation of large spatial
 ⁹⁸³ datasets. *Journal of Computational and Graphical Statistics*, 15(3):502–523, 2006. ISSN
 ⁹⁸⁴ 10618600. URL http://www.jstor.org/stable/27594195.

L. Gaultier and C. Ubelmann. Swot simulator documentation. Technical report, JPL, NASA, 2010.

I. J. Goodfellow, J. Pouget-Abadie, M. Mirza, B. Xu, D. Warde-Farley, S. Ozair, A. Courville, and
 Y. Bengio. Generative adversarial nets. In *Proceedings of the 27th International Conference on Neural Information Processing Systems - Volume 2*, NIPS'14, page 2672–2680, Cambridge,
 MA, USA, 2014. MIT Press.

A. Grigorievskiy, N. D. Lawrence, and S. Särkkä. Parallelizable sparse inverse formulation gaussian
 processes (spingp). 2017 IEEE 27th International Workshop on Machine Learning for Signal
 Processing (MLSP), pages 1–6, 2016.

J. Ho, A. Jain, and P. Abbeel. Denoising diffusion probabilistic models, 2020.

Y. Huang, Y. Zhang, and J. A. Chambers. A novel kullback–leibler divergence minimization-based
 adaptive student's t-filter. *IEEE Transactions on Signal Processing*, 67(20):5417–5432, 2019.
 doi: 10.1109/TSP.2019.2939079.

⁹⁹⁷ D. P. Kingma and M. Welling. Auto-encoding variational bayes, 2022.

⁹⁹⁸ K. Kurosawa and J. Poterjoy. A statistical hypothesis testing strategy for adaptively blending
 ⁹⁹⁹ particle filters and ensemble kalman filters for data assimilation. *Monthly Weather Review*, 151

- (1):105 125, 2023. doi: 10.1175/MWR-D-22-0108.1. URL https://journals.ametsoc. 1000 org/view/journals/mwre/151/1/MWR-D-22-0108.1.xml. 1001
- F. Le Guillou. Cartographie dynamique de la topographie de l'océan de surface par assimilation 1002 de données altimétriques. Theses, Université Grenoble Alpes [2020-....], May 2022. URL 1003 https://theses.hal.science/tel-03775828. 1004
- F. Le Guillou, L. Gaultier, M. Ballarotta, S. Metref, C. Ubelmann, E. Cosme, and M.-H. Rio. 1005 Regional mapping of energetic short mesoscale ocean dynamics from altimetry: performances 1006 from real observations. EGUsphere, 2023:1-17, 2023. doi: 10.5194/egusphere-2023-509. URL 1007 https://egusphere.copernicus.org/preprints/2023/egusphere-2023-509/.

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- H. Lewy, K. Friedrichs, and R. Courant. Über die partiellen differenzengleichungen der mathema-1009 tischen physik. Mathematische Annalen, 100:32-74, 1928. URL http://eudml.org/doc/ 1010 159283. 1011
- X. Li, B. Wu, X. Zhu, and H. Yang. Consecutively missing seismic data interpolation based on 1012 coordinate attention unet. IEEE Geoscience and Remote Sensing Letters, 19:1-5, 2022. doi: 1013 10.1109/LGRS.2021.3128511. 1014
- F. Lindgren, H. Rue, and J. Lindström. An explicit link between gaussian fields and gaussian 1015 markov random fields: the stochastic partial differential equation approach. Journal of the Royal 1016 Statistical Society: Series B (Statistical Methodology), 73(4):423–498, 2011. doi: 10.1111/j. 1017 1467-9868.2011.00777.x. URL https://rss.onlinelibrary.wiley.com/doi/abs/10. 1018 1111/j.1467-9868.2011.00777.x. 1019

S. A. Martin, G. E. Manucharyan, and P. Klein. Synthesizing sea surface temper-1020 ature and satellite altimetry observations using deep learning improves the accuracy 1021 and resolution of gridded sea surface height anomalies. Journal of Advances in 1022 Modeling Earth Systems, 15(5):e2022MS003589, 2023. doi: https://doi.org/10.1029/ 1023 2022MS003589. URL https://agupubs.onlinelibrary.wiley.com/doi/abs/10. 1024 1029/2022MS003589. e2022MS003589 2022MS003589. 1025

L. Menut, B. Bessagnet, R. Briant, A. Cholakian, F. Couvidat, S. Mailler, R. Pennel, G. Siour, 1026 P. Tuccella, S. Turquety, and M. Valari. The chimere v2020r1 online chemistry-transport model. 1027

- Geoscientific Model Development, 14(11):6781–6811, 2021. doi: 10.5194/gmd-14-6781-2021. URL https://gmd.copernicus.org/articles/14/6781/2021/.
- S. Metref, E. Cosme, F. Le Guillou, J. Le Sommer, J.-M. Brankart, and J. Verron. Wide swath altimetric satellite data assimilation with correlated-error reduction. *Frontiers in Marine Science*, 6:822, 2020. ISSN 2296-7745. doi: 10.3389/fmars.2019.00822. URL
 https://www.frontiersin.org/article/10.3389/fmars.2019.00822.
- J.-M. Molines. meom-configurations/NATL60-CJM165: NATL60 code used for CJM165 experiment, Mar. 2018. URL https://zenodo.org/record/1210116\#.XDpmc89Khp8.

T. Nakamura, K. Fukami, K. Hasegawa, Y. Nabae, and K. Fukagata. Convolutional neural network
 and long short-term memory based reduced order surrogate for minimal turbulent channel flow.
 Physics of Fluids, 33(2), Feb. 2021. ISSN 1070-6631. doi: 10.1063/5.0039845. Publisher
 Copyright: © 2021 Author(s).

- A. Paszke, S. Gross, S. Chintala, G. Chanan, E. Yang, Z. DeVito, Z. Lin, A. Desmaison, L. Antiga,
 and A. Lerer. Automatic differentiation in pytorch. 2017.
- M. Pereira, N. Desassis, and D. Allard. Geostatistics for large datasets on riemannian manifolds:
 A matrix-free approach. *Journal of Data Science*, 20(4):512–532, 2022. ISSN 1680-743X. doi:
 10.6339/22-JDS1075.
- H. S. Price, R. S. Varga, and J. E. Warren. Applications of oscillation matrices to diequations. J.
 Math. Phys, 45, 1966.
- F. Rozet and G. Louppe. Score-based data assimilation. In A. Oh, T. Naumann,
 A. Globerson, K. Saenko, M. Hardt, and S. Levine, editors, *Advances in Neural Information Processing Systems*, volume 36, pages 40521–40541. Curran Associates,
 Inc., 2023. URL https://proceedings.neurips.cc/paper_files/paper/2023/file/
 7f7fa581cc8a1970a4332920cdf87395-Paper-Conference.pdf.
- J. J. Ruiz, M. Pulido, and T. Miyoshi. Estimating model parameters with ensemble-based data assimilation: A review. *Journal of the Meteorological Society of Japan. Ser. II*, 91(2):79–99, 2013.

61

M. A. Sacco, J. J. Ruiz, M. Pulido, and P. Tandeo. Evaluation of machine learning techniques
 for forecast uncertainty quantification. *Quarterly Journal of the Royal Meteorological Society*,
 148(749):3470–3490, 2022. doi: https://doi.org/10.1002/qj.4362. URL https://rmets.
 onlinelibrary.wiley.com/doi/abs/10.1002/qj.4362.

M. A. Sacco, M. Pulido, J. J. Ruiz, and P. Tandeo. On-line machine-learning forecast un certainty estimation for sequential data assimilation. *Quarterly Journal of the Royal Me- teorological Society*, n/a(n/a), 2024. doi: https://doi.org/10.1002/qj.4743. URL https:
 //rmets.onlinelibrary.wiley.com/doi/abs/10.1002/qj.4743.

S. Särkka and J. Hartikainen. Infinite-dimensional kalman filtering approach to spatio-temporal gaussian process regression. In N. D. Lawrence and M. Girolami, editors, *Proceedings of the Fifteenth International Conference on Artificial Intelligence and Statistics*, volume 22 of

Proceedings of Machine Learning Research, pages 993–1001, La Palma, Canary Islands, 21–23

¹⁰⁶⁷ Apr 2012. PMLR. URL https://proceedings.mlr.press/v22/sarkka12.html.

S. Särkka, A. Solin, and J. Hartikainen. Spatiotemporal learning via infinite-dimensional bayesian
 filtering and smoothing: A look at gaussian process regression through kalman filtering. *IEEE Signal Processing Magazine*, 30(4):51–61, 2013. doi: 10.1109/MSP.2013.2246292.

M. Seeger, A. Hetzel, Z. Dai, E. Meissner, and N. D. Lawrence. Auto-differentiating linear algebra,
 2019.

- ¹⁰⁷³ Z. Shi, X. Xu, X. Liu, J. Chen, and M.-H. Yang. Video frame interpolation transformer, 2022.
 ¹⁰⁷⁴ URL https://arxiv.org/abs/2111.13817.
- F. Sigrist, H. R. Künsch, and W. A. Stahel. Stochastic partial differential equation based modelling
 of large space-time data sets. *Journal of the Royal Statistical Society: Series B (Statistical Methodology)*, 77(1):3–33, 2015. doi: https://doi.org/10.1111/rssb.12061. URL https://
 rss.onlinelibrary.wiley.com/doi/abs/10.1111/rssb.12061.
- H. Song, I. Hoteit, B. D. Cornuelle, X. Luo, and A. C. Subramanian. An adjoint-based adaptive ensemble kalman filter. *Monthly Weather Review*, 141(10):3343 3359, 2013. doi: 10.1175/MWR-D-12-00244.1. URL https://journals.ametsoc.org/view/journals/
 mwre/141/10/mwr-d-12-00244.1.xml.

- J. C. Strikwerda. *Finite Difference Schemes and Partial Differential Equations (Wadsworth & Brooks/Cole Mathematics Series)*. Wadsworth & BrooksCole Advanced Books & Software,
 Pacific Grove, 1989.
- Z. Su, J. Wang, P. Klein, A. F. Thompson, and D. Menemenlis. Ocean submesoscales as a key component of the global heat budget. *Nature Communications*, 9(1):775, Feb 2018.
 ISSN 2041-1723. doi: 10.1038/s41467-018-02983-w. URL https://doi.org/10.1038/s41467-018-02983-w.
- G. Taburet, A. Sanchez-Roman, M. Ballarotta, M.-I. Pujol, J.-F. Legeais, F. Fournier, Y. Faugere,
 and G. Dibarboure. DUACS DT2018: 25 years of reprocessed sea level altimetry products. 15
 (5):1207–1224, 2019. ISSN 1812-0784. doi: https://doi.org/10.5194/os-15-1207-2019. URL
 https://www.ocean-sci.net/15/1207/2019/. Publisher: Copernicus GmbH.
- P. Tandeo, P. Ailliot, J. Ruiz, A. Hannart, B. Chapron, A. Cuzol, V. Monbet, R. Easton, and
 R. Fablet. Combining Analog Method and Ensemble Data Assimilation: Application to the
 Lorenz-63 Chaotic System. In V. Lakshmanan, E. Gilleland, A. McGovern, and M. Tingley,
 editors, *Machine Learning and Data Mining Approaches to Climate Science*, pages 3–12.
 Springer, 2015.
- P. Tandeo, P. Ailliot, M. Bocquet, A. Carrassi, T. Miyoshi, M. Pulido, and Y. Zhen. A review of innovation-based methods to jointly estimate model and observation error covariance matrices in ensemble data assimilation. *Monthly Weather Review*, 148(10):3973 – 3994, 2020. doi: 10.1175/
 MWR-D-19-0240.1. URL https://journals.ametsoc.org/view/journals/mwre/148/
 10/mwrD190240.xml.
- J. Tong, L. Xie, W. Yang, and K. Zhang. Probabilistic decomposition transformer for time series forecasting, 2022.
- P.-Y. Traon. F. Nadal, and N. Ducet. An improved mapping method of mul-1106 tisatellite altimeter data. Journal of Atmospheric and Oceanic Technology, 15 1107 (2):522-534,1998. doi: 10.1175/1520-0426(1998)015<0522:AIMMOM>2.0.CO; 1108 URL https://journals.ametsoc.org/view/journals/atot/15/2/1520-0426\ 2. 1109 _1998_015_0522_aimmom_2_0_co_2.xml. 1110

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C. Ubelmann, P. Klein, and L.-L. Fu. Dynamic Interpolation of Sea Surface Height and Potential Applications for Future High-Resolution Altimetry Mapping. *Journal of Atmospheric and Oceanic Technology*, 32(1):177–184, Oct. 2014. ISSN 0739-0572. doi: 10. 1175/JTECH-D-14-00152.1. URL http://journals.ametsoc.org/doi/abs/10.1175/ JTECH-D-14-00152.1.

C. Ubelmann, P. Klein, and L.-L. Fu. Dynamic interpolation of sea surface height and potential applications for future high-resolution altimetry mapping. *Journal of Atmospheric and Oceanic Technology*, 32(1):177 – 184, 2015. doi: 10.1175/JTECH-D-14-00152.1. URL https://

journals.ametsoc.org/view/journals/atot/32/1/jtech-d-14-00152_1.xml.

C. Ubelmann, G. Dibarboure, L. Gaultier, A. Ponte, F. Ardhuin, M. Ballarotta, and Y. Faugère.
 Reconstructing ocean surface current combining altimetry and future spaceborne doppler data.
 Journal of Geophysical Research: Oceans, 126(3):e2020JC016560, 2021. doi: https://doi.org/
 10.1029/2020JC016560. URL https://agupubs.onlinelibrary.wiley.com/doi/abs/
 10.1029/2020JC016560. e2020JC016560 2020JC016560.

P. Virtanen, R. Gommers, T. E. Oliphant, M. Haberland, T. Reddy, D. Cournapeau, E. Burovski,
P. Peterson, W. Weckesser, J. Bright, S. J. van der Walt, M. Brett, J. Wilson, K. J. Millman,
N. Mayorov, A. R. J. Nelson, E. Jones, R. Kern, E. Larson, C. J. Carey, İ. Polat, Y. Feng,
E. W. Moore, J. VanderPlas, D. Laxalde, J. Perktold, R. Cimrman, I. Henriksen, E. A. Quintero,
C. R. Harris, A. M. Archibald, A. H. Ribeiro, F. Pedregosa, P. van Mulbregt, and SciPy 1.0
Contributors. SciPy 1.0: Fundamental Algorithms for Scientific Computing in Python. *Nature Methods*, 17:261–272, 2020. doi: 10.1038/s41592-019-0686-2.

H. Wackernagel. *Multivariate Geostatistics. An Introduction with Applications*. Springer-Verlag
Berlin Heidelberg, New York, 2003. doi: 10.1007/978-3-662-05294-5. ISBN 978-3-54044142-7.

P. Whittle. On stationary processes in the plane. *Biometrika*, 41(3-4):434–449, 1953. doi: 10.

1093/biomet/41.3-4.434. URL http://biomet.oxfordjournals.org/content/41/3-4/
434.short.

¹¹³⁸ X. Zhang, M. A. Nawaz, X. Zhao, and A. Curtis. Chapter two - an introduction to variational ¹¹³⁹ inference in geophysical inverse problems. In C. Schmelzbach, editor, *Inversion of Geophysical* Data, volume 62 of Advances in Geophysics, pages 73–140. Elsevier, 2021. doi: https://
 doi.org/10.1016/bs.agph.2021.06.003. URL https://www.sciencedirect.com/science/
 article/pii/S0065268721000030.

Y. Zhen, P. Tandeo, S. Leroux, S. Metref, T. Penduff, and J. L. Sommer. An adaptive optimal interpolation based on analog forecasting: Application to ssh in the gulf of mexico. *Journal of Atmospheric and Oceanic Technology*, 37(9):1697 – 1711, 2020. doi:
10.1175/JTECH-D-20-0001.1. URL https://journals.ametsoc.org/view/journals/
atot/37/9/jtechD200001.xml.